# **LOGISTICS and PROCESSING REPORT**

of the AIRBORNE MAGNETIC AND GEOTEM ELECTROMAGNETIC MULTICOIL SURVEY

in BERNALILLO, SANDOVAL and SANTA FE counties, NEW MEXICO

for the U.S. GEOLOGICAL SURVEY

Job  $N^{\circ}$  319 August, 1997 Ottawa, Ontario

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## **INTRODUCTION**

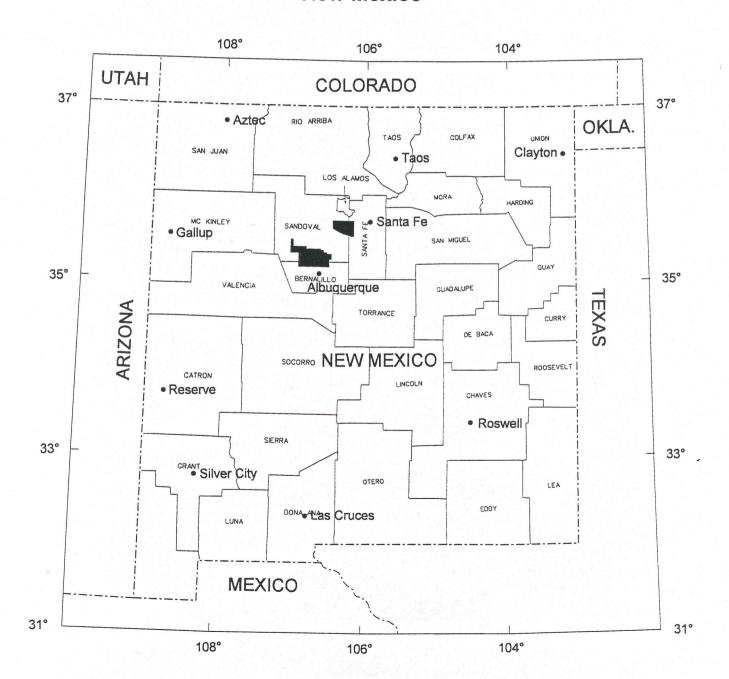
From February 18<sup>th</sup> to March 9<sup>th</sup>, 1997, an electromagnetic and magnetic survey was flown by Geoterrex-Dighem on behalf of the United States Geological Survey. Three areas were flown, namely the Cochiti Indian reservation, the Rio Rancho and the Rio Puerco blocks, in Bernalillo, Sandoval and Santa Fe counties, New Mexico. In all, 2338 line miles of data were collected.

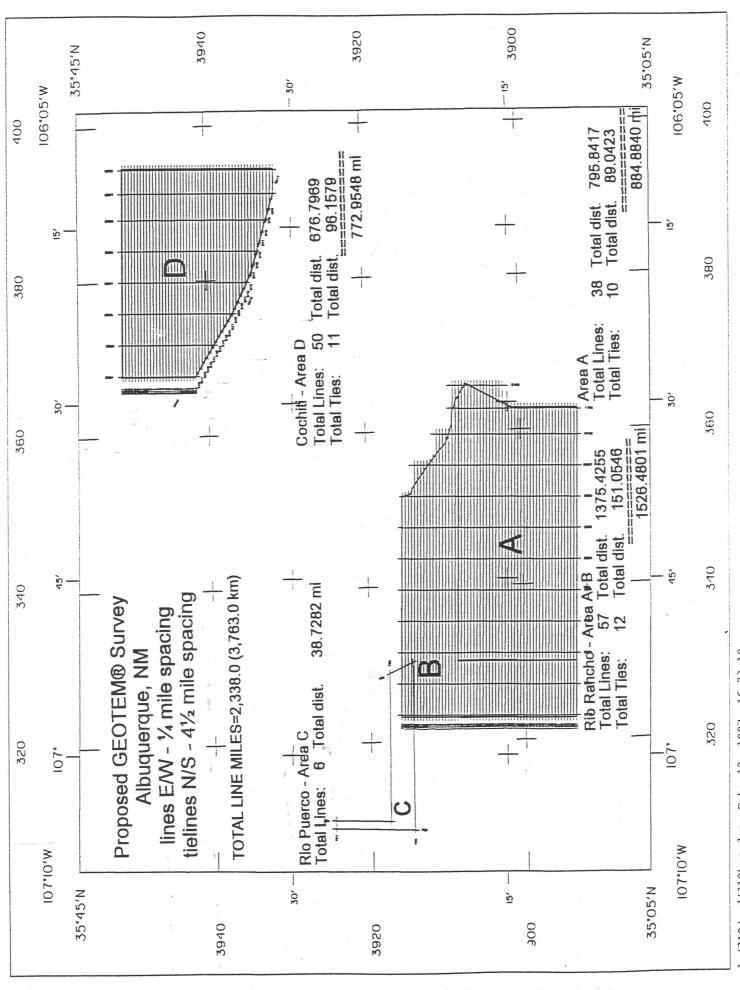
The survey data was compiled and processed in the Geoterrex-Dighem Ottawa office and is presented as maps of the total intensity magnetics with flight path, and the terrain clearance, multiparameter profiles, conductivity-depth-transform grids for each survey line and digital archive files.

# SURVEY LOCATION

# Cochiti Indian Reservation Rio Rancho area Rio Puerco reconnaissance lines

# **New Mexico**





1:/319/cad/319kms.dgn Feb. 12, 1997 16:23:10

#### **SURVEY OPERATIONS**

## **Location of the Survey Area**

The survey areas are in Bernalillo, Sandoval and Santa Fe counties, New Mexico. The Cochiti Indian reservation block is located west of Santa Fe; the Rio Rancho block is just north of Albuquerque; and the Rio Puerco test lines are located at the northwestern corner of the Rio Rancho block.

The base of operation was Albuquerque, New Mexico.

# Aircraft and Geophysical On-Board Equipment

Aircraft CASA C-212 twin turbo-prop.

Operator Geoterrex-Dighem

Registration C-FDPP

Survey Speed 125 knots/145 mph/65m/sec.

Magnetometer Scintrex Cs-2 single cell cesium vapour, towed-bird

installation, sensitivity =  $0.01 \text{ nT}^1$ , sampling rate = 0.1 sec., ambient range 20,000 to 100,000 nT. The general noise envelope was kept below 0.5 nT. Nominal sensor height of 73

metres above ground.

Electromagnetic system GEOTEM multicoil system

System parameters

*Transmitter.* vertical axis loop of 232 m<sup>2</sup>,

<sup>1</sup> One gamma is equivalent to the S.I. unit nanotesla (nT).

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number of turns: 3,

nominal height above ground of 120

metres.

Receiver. multicoil system (x, y and z) with a final

> recording rate of 4 samples/second, for the recording of 20 channels of x, y and z-coil data. Nominal height above ground of 70 metres, placed 125 m behind the centre of

the transmitter loop.

Base frequency: 30 Hz 4036 ì s Pulse width: Pulse delay: 130.2 ì s 12501 is Off-time: 130.2 is Point value:

Transmitter: Current of 935 amperes,

dipole moment of 6.5 x 10<sup>5</sup>Am<sup>2</sup>.

Window mean delay times in microseconds Receiver:

from the end of the pulse:

channel 1: 1106 channel 11: 9569 channel 2: 1432 channel 12: 11523 channel 3: 1822 channel 13: 846 channel 4: 2278 channel 14: 586 channel 5: 2864 channel 15: 390 channel 6: 3580 channel 16: 260 channel 7: 4427 channel 17: -455 channel 8: 5403 channel 18: -1757 channel 9: 6510 channel 19: -3059 channel 10: 7877 channel 20: -3841

Geoterrex-Dighem GEODAS.

RMS GR-33, showing the total magnetic field at 2 vertical scales, the radar and barometric altimeters, the time-constant filtered traces of the x-coil channels 1 to 12 and the on-time channel 20, the raw traces of the x and z-coil channel 1 and 12, the EM primary field, the powerline monitor, the 4<sup>th</sup> difference of the magnetics, the x and z-coil earth's field

monitors and the fiducials.

**Barometric Altimeter** Rosemount 1241M, sensitivity 1 foot, 1 sec. recording interval.

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**Digital Acquisition** 

Analog Recorder

Radar Altimeter TRT AHV-8, accuracy 2%, sensitivity one foot, range 0 to

2,500 feet, 1 sec. recording interval.

Camera Panasonic colour video, super VHS, model WV-CL302.

Electronic Navigation Sercel GPS receiver NR103, 1 sec. recording interval, with a

resolution of 0.00001 degree and an accuracy of ±10 m.

# **Base Station Equipment**

Magnetometer: Scintrex H8 single cell split beam cesium vapour, mounted in a

magnetically quiet area, measuring the total intensity of the earth's magnetic field in units of 0.01 nT at intervals of 1

second, within a noise envelope of 0.10 nT.

GPS Receiver: SERCEL NR103 V2.3, measuring all GPS channels, for up to

10 satellites.

Computer: Toshiba laptop, model T4600, 33 MHz, 486.
Converter: Picodas, model MEP710 3/10901 GTS 780008.

Printer: Kodak Diconix 150 plus.

Battery Backup.

# **Field Office Equipment**

Video Playback: Panasonic Super VHS with an 8" BRULE Colour Monitor.

Computer: Dell Latitude LM laptop with 800 MB hard drive.

Plotter: Design Mate 24" Calcomp with multiple colours and pen types,

quality - 2000 steps/inch.

Printer: Hewlett Packard Deskjet 660C.

DAT Tape Drive: Conner, 90 metre tapes (1300 MG).

Hard Drive: Seagate 4 gigabyte

## **Survey Specifications**

Altitude: The survey was flown at a mean terrain clearance of 120

metres.

The altitude tolerance was limited to  $\pm 15$  m, not to be exceeded over a distance greater than 3 km, unless required for safety or air regulations.

Traverse Lines: Spacing of 400 m, direction E-W.

The separation between flight lines was not to exceed 600 m over a distance greater than 4500 m.

**Control lines:** Spacing of 7250 m, direction N-S.

**Diurnal variation:** Acceptable variations were limited to 7.5 nT deviations from a

one minute chord.

**Noise levels:** The noise level on the magnetic data was not to exceed ±0.25

nT over a distance greater than 3 km.

The noise level on the electromagnetic data was not to exceed ±30 ppm over a distance greater than 3 km, as displayed on

the analogue traces of the late-time x-coil channels.

Lines or portions of lines flown during any of the above deviations were reflown. These reflights began and ended by crossing control lines and overlapping good data.

# **Survey Coverage**

In the Cochiti area, 773 line miles of data were collected.

In the Rio Rancho area, 1526 line miles of data were collected.

In the Rio Puerco area, only 39 line miles of data were collected. These were for general reconnaissance purposes.

Tests and Calibrations (flown from the Sierra Vista airport, March 14, 1997)

## System lag for the Magnetic and Electromagnetic response

A suite of lines were flown over a bridge, near the airport, along bearings 140° and 320°, and checked against the video positions to verify the system lag of the magnetic and EM response.

The average system lag from all passes were as follows:

Magnetics: 3.6 seconds Electromagnetics: 4.0 seconds

A sample of the response over the bridge is provided in figure 1.

# GPS Accuracy check (cloverleaf)

The accuracy of the GPS navigation system and its synchronization with the on-board video system was verified by flying a cloverleaf pattern at survey altitude over the basestation GPS antenna, set-up at the Sierra Vista airport (Easting of 563741, Northing of 3495176, datum Clarke 1866).

#### Control from the video:

<u>Pass</u>	<u>Dir</u>	<u>Altitude</u>	<u>Fiducial</u>	Video Position
1	W	380 ft	61181.55	3 m N of mark
2	Е	381 ft	61319.75	1 m N of mark
3	W	380 ft	61491.85	4 m N of mark
4	Е	383 ft	61624.85	1 m S of mark
5	S	383 ft	61841.35	8 m W of mark
6	Ν	431 ft	61959.98	7 m E of mark
7	S	392 ft	62171.50	5 m W of mark
8	Ν	406 ft	62329.35	5 m E of mark

# Results after post flight differential corrections applied:

Scatter of solutions inside an error box of 15 x 20 metres.

Offset of mean position from all passes relative to the mark is less than one metre.

Synchronization difference of the GPS position and the video image is 0.5 second.

Figure 2: Sketch of the calibration site

**Figure 3**: Trajectory of the aircraft over the calibration site

## Altimeter Calibration

A series of altitude passes were flown over the runway (elevation of 4612 feet above sea level), ranging from 300 to 800 feet terrain clearance, as monitored by the on-board radar altimeter.

**Figure 4** shows the relationship of the barometric elevation (as measured by the GPS receiver) and the radar altitude (adjusted for the runway elevation). The plot shows a good linear relationship with a correlation factor of 0.913. The increased spreading of points with the increase in altitude reflects the apparent shift between the radar and barometric altimeters when the aircraft is banked and turning. The barometric information displays true information,

while the radar indicates an exaggerated height since the beam is no longer focused directly beneath the aircraft, but at an angle away from it.

**Figure 5** explains this discrepancy between the radar and barometric altitudes more clearly. The "peaks" represent the aircraft in turnarounds, with the radar trace displaying a greater deviation; the "troughs" represent the aircraft returning to straight and level flying over the runway between turnarounds. The areas of straight and level flying (troughs) are the only regions where the direct comparison of the radar and barometric elevation is valid.

#### **Field Crew**

The base of field operations was Albuquerque, New Mexico.

Pilots: G. Stonehouse, G. Mueller, D. Wiens, M. Mellett

Electronics Operators: D. Patzer, R. Penton, J. Moore

Engineer: R. Constapel Geophysicist: R. Williams

#### **Production Statistics**

Flying was done between February 18<sup>th</sup> and March 9<sup>th</sup>, 1997.

Total production: 3762 km

Number of production flights: 16

Hours of production flying: 42.7 hours

Number of km/hour of production flying: 88.1 km/hr. Number of km/average production flight: 235 km Number of hours/average production flight: 2.7 hours

Number of days lost to:

Equipment: 1 Weather: 4.5 Testing: 0.5

Pilot training and/or rest: 5

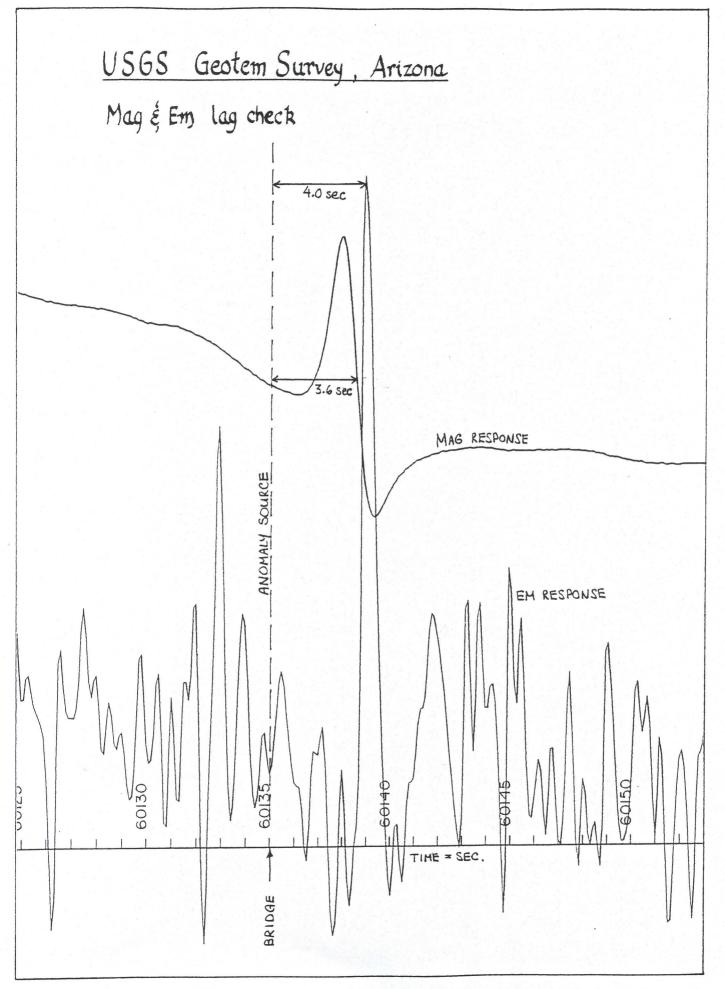
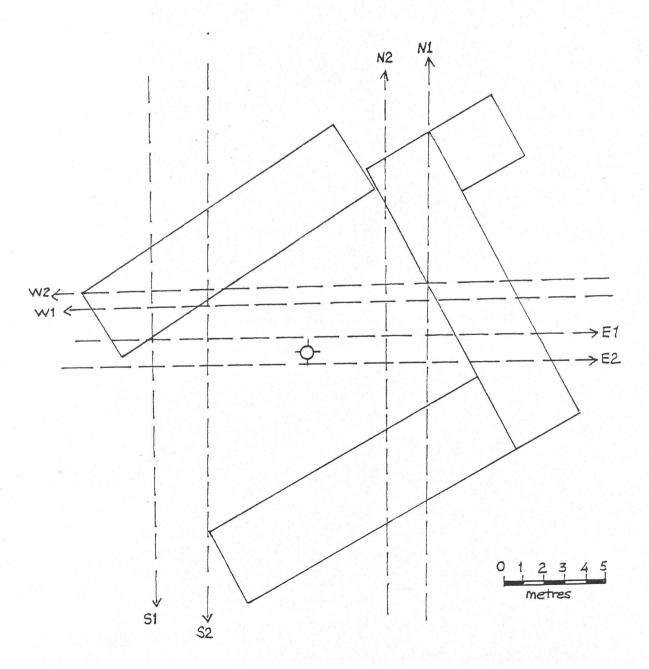


FIGURE 1

# U.S.G.S Geotem Survey, Arizona

GPS A CCURACY CHECK Sierra Vista Airport

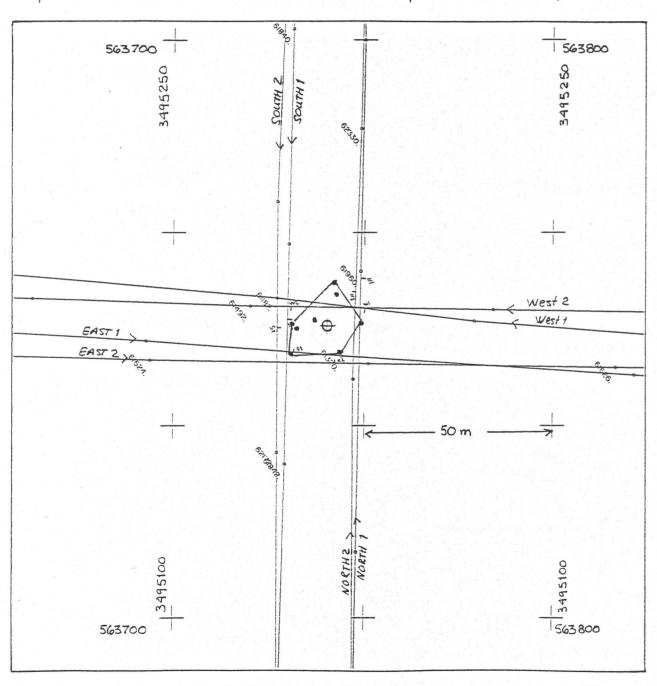
Location of mark (Clarke 1866) UTM-X = 563741 metres UTM-Y = 3495176 metres



# USGS Geotem Survey, Arizona

GPS ACCURACY CHECK Sierra Vista Airport

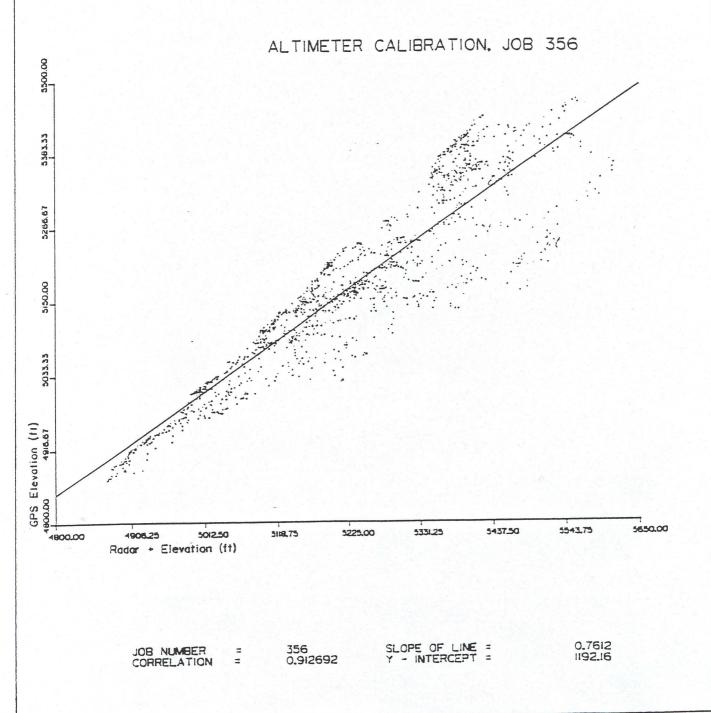
= Location of reference mark . = Video position from each pass

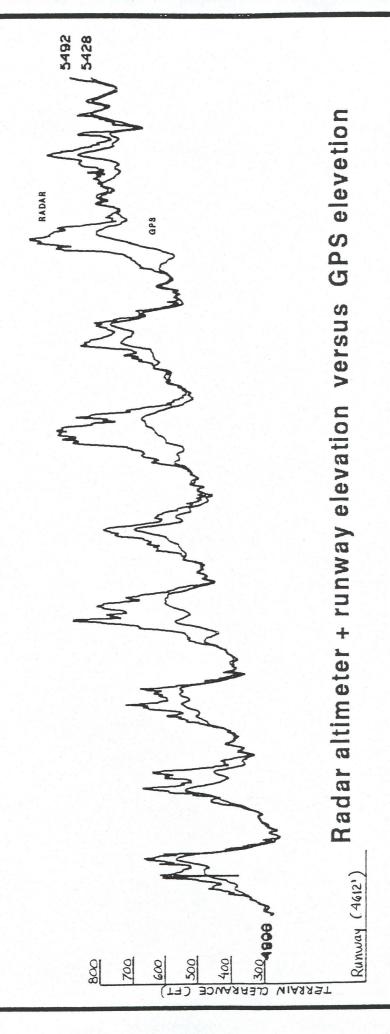


# USGS Geotem Survey, Arizona

Altimeter Calibration Sierra Vista Airport

Runway elevation = 4612 feet







## **QUALITY CONTROL AND COMPILATION PROCEDURES**

#### In the Field

After each flight, all analog records were examined as a preliminary assessment of the noise level of the recorded data. Altimeter deviations from the prescribed flying altitudes were also closely examined as well as the general condition of the diurnal activity, as recorded on the base stations.

All digital data were verified for validity and continuity. The data from the aircraft and base station was transferred to the PC's hard disk. Basic statistics were generated for each parameter recorded; these included the minimum, maximum and mean values, the standard deviation and any null values located. Editing of all recorded parameters for spikes or datum shifts was done, followed by final data verification via an interactive graphics screen with onscreen editing and interpolation routines. Any of the recorded parameters could be plotted back at a suitable scale on the field pen plotter.

The quality of the navigation was controlled on a daily basis by recovering the flight path of the aircraft. The Trajecto correction procedure employs the raw ranges from the base station to create improved models of clock error, atmospheric error, satellite orbit, and selective availability. These models are used to improve the conversion of aircraft raw ranges to aircraft position. The Trajecto-corrected GPS was plotted back daily in the field on the pen plotter and checked for speed busts.

Checking all data for adherence to specifications was carried out in the field by the Geoterrex-Dighem field operations manager/data technician.

More details on the individual processing steps applied to each parameter are given in Appendix A of this report.

# IV

## DATA PROCESSING IN THE OFFICE

# Flight Path Recovery

Data used: GPS positions recalculated from the recorded raw range data, differentially corrected and converted to UTM metres.

# Final positions:

Projection: Universal Transverse Mercator

Central meridian: 109°W

False Easting: 500,000 metres

False Northing: 0 metres Scale factor: 0.9996 Spheroid: Clarke 1866

## Altitude Data (Radar and GPS)

Noise editing: Alfatrim median filter used to eliminate the 2 high and 2 low values from the statistical distribution of a 9 point sample window.

Noise filtering: Triangular filters set to remove radar wavelengths less than 6 seconds and amplitudes below 25 feet; and GPS wavelengths up to 4 seconds and amplitudes up to 2.5 metres.

*Gridding:* A grid of the terrain clearance was produced from the radar altimeter values, as a general reference for evaluation of the survey results.

# **Diurnal magnetics**

Noise editing: Alfatrim median filter used to eliminate the 2 high and 2 low values from the statistical distribution of a 9 point sample window.

Culture editing: Polynomial interpolation via a graphic screen editor.

Noise filtering: Triangular filter set to remove wavelengths of less than 4 seconds and amplitudes up to 0.5 nT.

Extraction of long wavelength component: Low pass filter set to retain only wavelengths of greater than 50 seconds.

## **Magnetics**

Lag correction: 3.6 seconds

*Noise editing:* 4<sup>th</sup> difference editing routine set to remove spikes of greater than 0.5 nT, followed by an alfatrim median filter eliminating 2 high and 2 low value from its calculation over a 9 point window.

*Noise filtering:* Triangular filter set to remove noise events having a wavelength of less than 0.5 second and an amplitude of less than 0.3 nT.

*Diurnal subtraction:* The long wavelength component of the diurnal (greater than 50 seconds) was removed from the data, prior to the levelling analysis.

Levelling: The levelling consisted of applying a micro-levelling routine to the gridded data to remove small residual effects. The difference in the gridded datasets before and after the application of the micro-levelling routine were computed and extracted along the original survey lines to be stored in the final line dataset as the final magnetic compensation values.

*Gridding:* The data was gridded with a 25 m grid cell size, using a minimum curvature routine. *Regional:* The regional values were computed from the 1995 coefficients for the date of 1997.16 at an elevation of 1900 m a.s.l.

#### **Electromagnetics**

Lag correction: 4.0 seconds

Data correction: The x, y and z-coil data were processed from the 20 raw channels recorded at 4 samples per second.

The EM data measured by the three coils, x, y and z, is collected in  $pv/m^2$  which is an unnormalized value permitting a direct comparison of the signal between the three coil sets. The conversion of the data from  $pv/m^2$  to ppm is a function of the effective area of each coil and the voltage of the primary signal received by each coil. The effective area of the coil is identical for all coils, but the value of the primary field voltage differs for each coil due to coupling effects such that the derived ppm value becomes a relative value for a particular coil set only. If quantitative analysis of the data is required, where the response from the different coil sets are to be referenced to each other, then the data should be left in unnormalized  $pv/m^2$  units which is the true representation of the signal measured.

For the current configuration of 30 Hz with a 4 ms pulse, the correspondence between pv/m<sup>2</sup> and ppm is roughly:

28:1 for the x-coil; 16:1 for the z-coil; and 2:1 for the y-coil.

The final data is archived in pv/m<sup>2</sup> but the conversion factors to normalized the data back to ppm are given in a table in Appendix I. The conversion factors vary slightly from one flight to the next reflecting small variations in the received primary field from day to day.

The following processing steps were applied to the channel data from all coil sets:

- a) the data from channels 1 to 16 (off-time) and 20 (on-time) were corrected for drift in flight form (prior to cutting the recorded data back to the correct line limits) by passing a low order polynomial function through the baseline minima along each channel, via a graphic screen display;
- b) the data were edited for residual sferic spikes by examining the decay pattern of each individual EM transient. Bad decays (i.e. not fitting a normal exponential function) were deleted and replaced by interpolation;
- c) for the x and z-coils only, the data were initially corrected for incoherent, nondecaying low frequency noise events by analyzing the decay information through decay constant calculations, in order to separate the true signal from the low-frequency noise.
  - An adaptive filter was then applied to the data. This filter responds to local changes in the gradient in order to select and apply an appropriate time domain convolution ranging from very narrow to wide, depending on the local character of the anomaly. The suite of filters selected were aimed at noise frequencies ranging from 2 Hz to 0.18 Hz;
- d) for the y-coil data, spikes were removed using an alfatrim median filter to isolate extreme high and low values. This was followed by the application of a small running average filter.
  - An adaptive filter was then applied to the data. This filter responds to local changes in the gradient in order to select and apply an appropriate time domain convolution ranging from very narrow to wide, depending on the local character of the anomaly;
- e) the filtered data from the x, y and z-coils were then re-sampled to a sample rate of 5 samples/sec and combined into a common file for archiving.

Decay constant calculation: An off-time decay constant was calculated by fitting the channel information to an exponential function. The decay constant was calculated using channels 2 to 12 from the z-coil data (mean delay times of 1432 to 11,523 i sec). This parameter was provided in preliminary form only for the purpose of assessing the general EM response over the survey blocks.

Conductivity-Depth-Transform (CDT): Conductivity-Depth-Transforms were computed along each survey line, resulting in a vertical pseudo-section of resistivity values. In this calculation, the complete waveform (all 20 channels) is used, fitting the combined data from the x and z-coils to a layered earth model. The derived resistivity grid values are then adjusted such that the top of the grid reflects the true terrain topography.

For the **Cochiti block**, the maximum depth of investigation was set at 1300 m from a reference barometric elevation of 2500 m.

For the **Rio Rancho block**, the maximum depth of investigation was set at 1100 m from a reference barometric elevation of 2200 m.

For the **Rio Puerco reconnaissance lines**, the maximum depth of investigation was set at 1000 m from a reference barometric elevation of 2300 m.

Further description on the content of the CDT grids and their format is given in Appendix G.

Information on the CDT algorithm itself is given in the article published by P. Wolfgram and G. Karlik, in Exploration Geophysics, 1995, vol. 26, pp. 179-185.

Additional Reference grids: The following data grids were also generated for each survey block as a general reference when evaluating the dataset:

- amplitude of the x-coil channel 2;
- amplitude of the x-coil channel 6;
- amplitude of the x-coil channel 10;
- amplitude of the z-coil channel 2;
- amplitude of the z-coil channel 6;
- amplitude of the z-coil channel 10.

# V

## **FINAL PRODUCTS**

# **Digital Archives**

The line data are archived in 3 files having standard ASCII format. The 1<sup>st</sup> file contains a flight line identification, eastings and northings and the 20 channels of processed x, y and z-coil data. The 2<sup>nd</sup> file is the same as the 1<sup>st</sup> except having the raw data. The 3<sup>rd</sup> file contains all the magnetic, altimeter and other secondary data as described in Appendix F (Archive Format Description).

The gridded data are archived in GEOSOFT ASCII format as a .GXF file. There are grids of channels 2, 6, 10, x and z-coil, magnetics, radar altimeter and the individual CDT grids for each survey line.

#### **Profile Data**

1 set of multiparameter profiles displaying the following information at a horizontal scale of 1:24,000, on paper.

- Radar altimeter
- EM Primary field
- Residual magnetics
- 12 off-time x-coil channels
- 12 off-time y-coil channels
- 12 off-time z-coil channels
- Hz monitor

# Maps

1 set of maps at 1:100,000 scale, presented on a clear UTM base with flight path.

- Total Field magnetic contours; delivered as 3 mylars
- Terrain Clearance contours; delivered as 1 mylar

Note: Also delivered were the digital reference EM waveform (parameter table) for each flight (a sample is included in Appendix H) and the digital data from the pre-survey tests and calibration in GEOSOFT ASCII format.

# Appendix A Field Processing of the Data

# Flight Path Recovery

- 1. Recover GPS data from base station.
- Recover GPS data from aircraft file.
- 3. Combine base station and aircraft GPS data to a common file.
- 4. Apply differential corrections and compute Lat/Long fixes from the raw range data (TRAJECTOGRAPHY); GPS lag of 0.5 sec used.
- 5. Change spheroid from WGS84 to Clarke 1866 using the following offsets:

Dx = 8 m

Dy = -163 m

Dz = -180 m

6. Convert Lat/Long fixes to UTM metres, using the following:

Central meridian: 105°W False Easting: 500,000 False Northing: 0

Scale factor: 0.9996 Ellipsoid: Clarke 1866

7. Compute speed from resulting positions to check for irregularities, time gaps, etc.

## Correction of Altimeter Data

- 1. Recover altimeter data from aircraft file.
- 2. Convert radar and GPS altitude data to feet and cm respectively.
- Remove spikes from the radar and GPS data using an alfatrim median filter to isolate the 2 extreme values (high or low) of a statistical distribution based on a window of 9 points.
- 4. Apply triangular convolutions to the altimeter data; aimed at wavelengths of 4 seconds or less and amplitudes of 2.5 m, or less for the GPS and wavelengths of 6 seconds or less and amplitudes of 25 feet or less for the radar.

# Processing of Magnetic Data from the Base Station

- 1. Recover magnetic data from the base station.
- 2. Edit the ground data for spikes, isolating the 2 extreme values out of a 9 point distribution (alfatrim median filter).
- 3. Remove culture events (when applicable) by polynomial interpolation via a graphic display/editing routine.
- 4. Apply a triangular smoothing convolution, aimed at removing wavelengths of 4 seconds or less with amplitudes of 0.5 nT or less.
- 5. Extract long wavelength component of the data (greater than 25 seconds) with a low pass filter, to be subtracted from the air data.

# Processing of Magnetic Data from the Aircraft

- 1. Recover data at full 10 samples/sec.
- 2. Adjust readings for a system lag of 3.6 seconds (36 samples).
- Edit data for spikes using a fourth difference routine, set for spikes of greater than 0.5 nT.
- 4. Apply triangular smoothing convolution, aimed at removing wavelength of less than 0.5 seconds with amplitudes less than 0.3 nT.
- 5. Remove the long wavelength component of the diurnal data (greater than 25 seconds).
- 6. Add back the mean value of the diurnal field.
- 7. Grid the resulting magnetic values, image and check for irregularities.

# Processing of GEOTEM Multicoil Data

- 1. Recover x, y and z-coil data in full flight format; x and z-coil channels at 4 samples/second.
- 2. Adjust readings for a system lag of 4.0 seconds (16 samples).
- 3. Review baseline positions for every channel of x and z-coil data via a graphic screen display and remove any visible drift component when required by graphically reconstructing and subtracting a low order polynomial. This procedure is only followed in resistive terrains when background "no response" sections along lines can be observed.
- 4. Isolate and remove, by interpolation, sferic events which are located in the data by analyzing the decay pattern at each sample point. All bad samples identified are labelled in the flag field. A decay threshold of 400 ppm was used, with the correction of no more than 2 consecutive samples allowed.
- 5. Remove any remaining small spikes in the data (x and z-coil channels) by using an alfatrim median filter to isolate the extreme value (high or low) of a statistical distribution based on a window of 5 sample points.
- 6. Correct the data for low frequency, incoherent noise elements (that do not correlate from channel to channel) in the data, by analyzing the decay patterns of channels 9 to 20 (OMEGA process). Decay threshold of 180 ppm used and the RMS error in fitting stored.
- 7. Noise filtering done using an adaptive filter technique based on time domain triangular

operators. Using a 2nd difference value to identify changes in gradient along each channel, minimal filtering (3 point convolution) is applied over the peaks of the anomalies, ranging in set increments up to a maximum amount of filtering in the resistive background areas (29 points for the z-coil data and 31 points for the x-coil data).

#### Creation of Final Files

- 1. Processed positioning, altimeter, magnetic and GEOTEM x, y and z-coil data is cutback from flight form to survey line limits, re-sampled and combined into a common file at 5 samples/sec.
- EM anomalies are automatically selected (using channel 6 as reference) and fitted to the vertical plate model to derive the conductivity-thickness product and the depth to source.
- 3. The automatic anomaly selection is reviewed and updated interactively via a graphic screen display/editing routine.

#### Field Products

- 1. Decay Constant from z-coil at 1:100,000 scale, deherringboned.
- 2. Preliminary total magnetic field contour map with flight path at 1:100,000 scale.

# Appendix B GEOTEM® EM System

# GEOTEM® ELECTROMAGNETIC SYSTEM

#### General

The operation of a towed-bird time-domain electromagnetic system (EM) involves the measurement of decaying secondary electromagnetic fields induced in the ground by a series of short current pulses generated from an aircraft-mounted transmitter. Variations in the decay characteristics of the secondary field (sampled and displayed as channels) are analyzed and interpreted to provide information about the subsurface geology. The response of such a system utilising a vertical-axis transmitter dipole and a horizontal-axis receiver coil has been documented by various authors including Palacky and West (1973, Geophysics, v. 38, p. 1145-1158).

The principle of sampling the induced secondary field in the absence of the primary field (during the "off-time") gives rise to an excellent signal-to-noise ratio and an increased depth of penetration compared to conventional continuous wave (frequency domain) electromagnetic systems. Such a system is also relatively free of noise due to air turbulence.

Through free-air model studies using the University of Toronto's Plate and Layered Earth programs it may be shown that the "depth of investigation" depends upon the geometry of the target. Typical depth limits would be 250 m below surface for a homogeneous half-space, 350 m for an inductively thin sheet or 200 m for a large vertical-plate conductor. These depth estimates are based on the assumptions that a significant response is recorded 1.4 ms after the primary field is turned off and that the overlying or surrounding material is resistive. If a preselected range of channels is deemed adequate to detect or resolve a given target, then the depth of investigation increases significantly.

In addition to substantial penetration, time-domain systems respond to a wide range of conductors. With measurements taken during off-time, significant effects are seen for values of 0.8 s or greater, thus responding to the majority of geological features deemed relevant in most exploration projects and many mapping applications.

The method also offers very good discrimination of conductor geometry. This ability to distinguish between flat-lying and vertical conductors combined with excellent depth penetration results in good differentiation of bedrock conductors from surficial conductors.

Most of the preceding discussion concerns theoretical responses for all time-domain installations of similar geometry. The factor which distinguishes the various systems is their ability to faithfully and completely represent the secondary field response.

# **Equipment Procedure**

GEOTEM® (GEOterrex Transient ElectroMagnetic system) is a time-domain towed-bird electromagnetic system incorporating a high-speed digital EM receiver. The primary electromagnetic pulses are created by a series of discontinuous sinusoidal current pulses fed into a three-turn transmitting loop surrounding the aircraft and fixed to the nose, tail and wing tips. The base frequency rate is selectable: 75, 90, 125, 150, 225 and 270 Hz. The standard system configuration (150 Hz, equivalent to 300 bipolar pulses per second) delivers a half-sine current pulse lasting 1030 µs followed by 2303 µs off-time. Present peak current through the loop is 600 A resulting in a primary magnetic dipole moment of 4.5 x 10<sup>5</sup>A.m². Optionally, the system can be reconfigured to generate a half-sine pulse lasting up to 1700 µs.

The receiver is an induction coil with a ferrite core. The receiving coil axis is aligned in the flight direction and mounted horizontally in a bird which is towed by the aircraft on a 135-metre cable. A non-magnetic tow cable is used to reduce noise levels. Mean terrain clearance for the aircraft is

120 m with the EM bird being situated nominally 56 m below and 123 m behind the aircraft.

For each primary pulse a secondary magnetic field is produced by decaying eddy currents in the ground. These in turn induce a voltage in the receiver coil which provides a measure of the electromagnetic field.

The measured signals pass through anti-aliasing filters and are then digitized with an analog-to-digital (A/D) converter at sampling rates of up to 40 kHz. The digital data flows from the A/D converter into an array processor where all the numerically intensive processing tasks, such as Fourier analysis, are carried out. The array processor is under the control of a multi-tasking minicomputer which provides all of the software management.

Operations which are carried out in the receiver are:

- 1. Compensation: During flight, the transmitter creates eddy currents within the structure of the aircraft that have a measurable effect at the bird. Compensation for this signal is implemented numerically within the receiver by a statistical analysis of the signal detected at the bird in absence of ground response (i.e., achieved by flying at an altitude such that no ground response is detectable). The observed signal is used to define a compensation signal which is removed from the observed signal to produce a null and thus effectively buck out any response due to the aircraft.
- 2. Normalization: All EM response channels are automatically calibrated and reduced to parts per million (ppm) of the primary field in the receiver. This is achieved by dividing the measured voltage by the voltage induced by the primary field at the bird.
- 3. Transient Analysis: Harmonic analysis permits the separation of specific types of noise from the signal in real time.
- 4. Digital Stacking: Stacking is carried out to reduce the effect of broadband noise on the data. Stacking time is 120 ms in areas where 50 Hz power lines are used and 133 ms where they are 60 Hz.
- 5. Windowing of data: The GEOTEM® digital receiver samples the secondary and primary electromagnetic field at 128 points per EM pulse and windows the signal in up to 20 time gates whose centres and widths are software selectable and which may be placed anywhere within or outside the transmitter pulse. This flexibility offers the advantage of arranging the gates to suit the goals of a particular survey, ensuring that the signal is appropriately sampled through its entire dynamic range.
- 6. Power Line Filtering: Digital comb filters are applied to the data during real-time processing to remove power line interference while leaving the EM signal undisturbed. The RMS power line voltage (at all harmonics in the receiver passband) are computed, displayed and recorded for each data stack.
- 7. Primary Field: The primary field at the towed sensor is measured for each stack and recorded as a separate data channel to assess the variation in coupling between the aircraft and the towed sensor induced by changes in system geometry.
- 8. Earth Field Monitor: A monitor of sensor coil motion noise induced by coil motion in the Earth's magnetic field is also extracted in the course of the real-time digital processing. This information is also displayed on the real-time chart as well as being recorded for post-survey diagnostic processes.

9. Noise/Performance: A monitor computes the RMS signal level on channel 3 over a running 10-second window. This monitor provides a measure of noise levels in areas of low ground response. This information is printed at regular intervals on the side of the flight record and is recorded for every data stack.

One of the major roles of the GEOTEM® digital receiver is to provide diagnostic information on system functions and to allow for identification of noise events, such as sferics, which may be selectively removed from the EM signal.

The receiver automatically calibrates its received signal with reference to the primary field in ppm and hence compensates for the transmitter drift. Due to the fact that the receiver is digital, receiver drift is minimal. These factors result in more reliable resistivity mapping where base level shifts can dramatically affect results.

GEOTEM®s high digital sampling rate yields maximum resolution of the secondary field. The absence of an analog system time-constant filter results in minimal signal distortion and, therefore, superior representation of the anomaly amplitudes and shapes.

# System Hardware

The GEOTEM® system is an integrated whole, consisting of the CASA 212 aircraft, the on-board hardware, and the software packages controlling the hardware.

The software packages in the GEODAS data acquisition system and in the HP1000-A600 GEOTEM® computer were developed in-house by Geoterrex. Likewise, certain elements of the hardware (GEOTEM® transmitter, system timing clock, towed-bird receiver system) were developed in-house.

Computer: Hewlett Packard HP1000-A600

The HP1000 is a rack-mounted minicomputer capable of one million instructions per second. It acts as a controlling host for the analogic array processor. Communication between the HP1000 and array processor is via DMA transfers. The software running on the HP1000 controls the actions of the array processor, and passes the resulting output data to the GEODAS for display and recording.

Analogic Array Processor: AP504

The array processor performs the real-time processing on the EM data. It is controlled by the HP1000. Receiver data comes to the AP504 from the ADC-510-60 analog-to-digital converter.

Analogic Analog-to-Digital Converter: ADC-510-60

The ADC-510-60 digitizes the receiver coil signal. It has 16-bit resolution and can digitize 2 differential input channels at sampling rates up to 52.5 kHz.

Transmitter System

The transmitter system drives high-current pulses of an appropriate shape and duration through the coils mounted on the CASA aircraft.

System Timing Clock

This subsystem provides appropriate timing signals to the transmitter, and also to the analog-to-digital converter, in order to produce output pulses and capture the ground response.

Towed-Bird Receiver System

A single coil is mounted inside a towed bird which typically is 56 metres below and 123 metres behind the aircraft. The coil axis is parallel to the longitudinal axis of the aircraft. (A second bird, housing the magnetometer sensor, is typically 47 metres below and 82 metres behind the aircraft.)

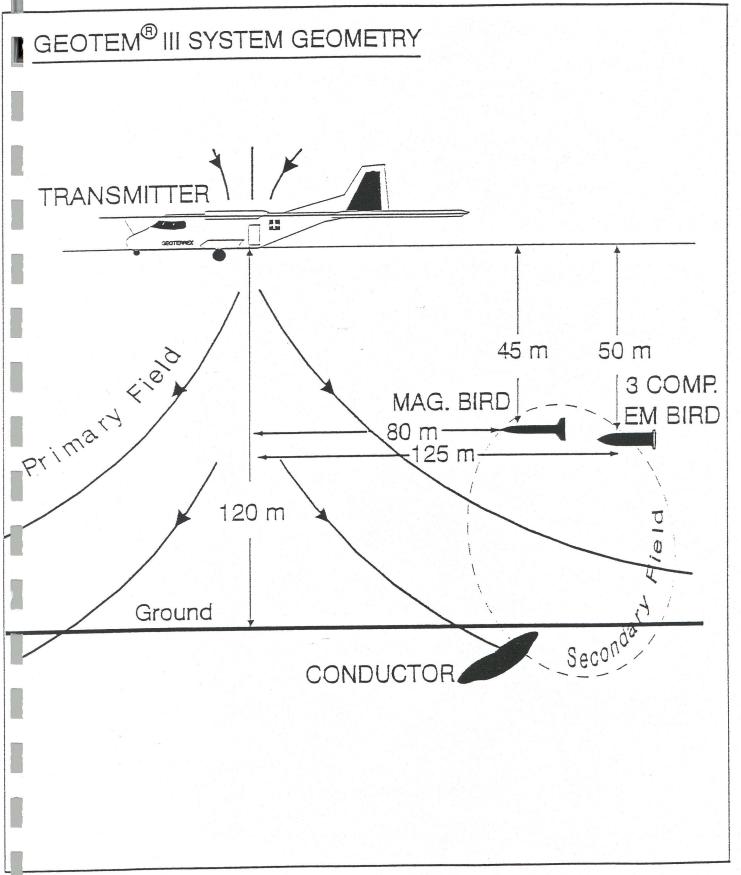
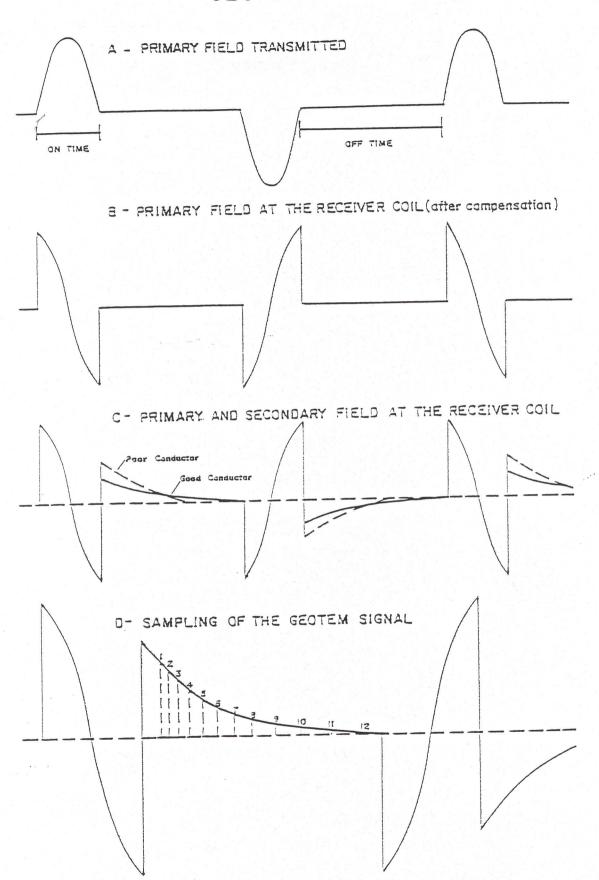


Figure 1

# GEOTEM® SIGNAL



# Appendix C GEOTEM® Interpretation Notes

# GEOTEMAT INTERPRETATION

# I. INTRODUCTION

The basis of the transient electromagnetic (EM) geophysical surveying technique relies on the premise that changes in the primary EM field produced in the transmitting loop will result in eddy currents being generated in any conductors in the ground. The eddy currents then decay to produce a secondary EM field which may be sensed as a voltage in the receiver coil.

GEOTEM (GEOterrex Transient ElectroMagnetic system) is an airborne transient (or time-domain) towed-bird EM system incorporating a high-speed digital receiver which records the secondary field response with a high degree of accuracy. Most often the total magnetic field is recorded concurrently.

Although the approach to GEOTEM interpretation varies from one survey to another depending on the type of data presentation, objectives and local conditions, the following generalizations may provide the reader with some helpful background information.

The main purpose of the interpretation is to determine the probable origin of the conductors detected during the survey and to suggest recommendations for further exploration. This is possible through an objective analysis of all characteristics of the different types of conductors and associated magnetic anomalies, if any. If possible the airborne results are compared to other available data. A certitude is seldom reached, but a high probability is achieved in identifying the conductive causes in most cases. One of the most difficult problems is usually the differentiation between surface conductors and bedrock conductors.

# II. TYPES OF CONDUCTORS

# A. Bedrock Conductors

The different types of bedrock conductors normally encountered are the following:

Graphites. Graphitic horizons (including a large variety of carbonaceous rocks) occur in sedimentary formations of the Precambrian as well as in volcanic tuffs, often concentrated in shear zones. They correspond generally to long, multiple conductors lying in parallel bands. They have no magnetic expression unless associated with pyrrhotite or magnetite. Their conductivity is variable but generally high.

GEOTEM<sup>a</sup>: Registered Trade Mark of Geoterrex Limited.

2. <u>Massive sulphides.</u> Massive sulphide deposits usually manifest themselves as short conductors of high conductivity, often with a coincident magnetic anomaly. Some massive sulphides, however, are not magnetic, others are not very conductive (discontinuous mineralization), and some may be located among formational conductors so that one must not be too rigid in applying the selection criteria.

In addition, there are syngenetic sulphides whose conductive pattern may be similar to that of graphitic horizons but these are generally not as prevalent as graphites.

- 3. <u>Magnetite and some serpentinized ultrabasics.</u> These rocks are conductive and very magnetic.
- 4. <u>Manganese oxides.</u> This mineralization may give rise to a weak EM response.

# B. Surficial Conductors

- 1. Beds of clay and alluvium, some swamps, and brackish ground water are usually poorly conductive to moderately conductive.
- 2. Lateritic formations, residual soils and the weathered layer of the bedrock may cause surface anomalous zones, the conductivity of which is generally low to medium but can occasionally be high. Their presence is often related to the underlying bedrock.

# C. <u>Cultural Conductors (Man-Made)</u>

- 1. <u>Power lines.</u> These frequently, but not always, produce a conductive type of response on the GEOTEM record. In the case of direct radiation of its field, a power line is easily recognized by a GEOTEM anomaly which exhibits phase changes between different channels. In the case of a grounded wire, or steel pylon, the anomaly may look very much like a bedrock conductor.
- 2. Grounded fences or pipelines. These will invariably produce responses much like a bedrock conductor. Whenever they cannot be identified positively, a ground check is recommended.
- 3. General culture. Other localized sources such as certain buildings, bridges, irrigation systems, tailings ponds etc., may produce GEOTEM anomalies. Their instances, however, are rare and often they can be identified on the visual path recovery system.

# III. ANALYSIS OF THE CONDUCTORS

The apparent conductivity alone is not generally a decisive criterion in the analysis of a conductor. In particular, one should note:

- its shape and size,
- all local variations of characteristics within a conductive zone,
- any associated geophysical parameter (e.g. magnetics),
- the geological environment,
- the structural context, and
- the pattern of surrounding conductors.

The first objective of the interpretation is to classify each conductive zone according to one of the three categories which best defines its probable origin. The categories are cultural, surficial and bedrock. A second objective is to assign to each zone a priority rating as to its potential as an economic prospect.

### A. Cultural Conductors

The majority of cultural anomalies occur along roads and are accompanied by a response on the power line monitor. (This monitor is set to 50 or 60 Hz, depending on the local power grid.) Power lines are the most common source of the anomalies and many are recognized immediately by virtue of phase reversals or an abnormal rate of decay. A certain number yield normal GEOTEM anomalies which could be mistaken for bedrock responses. There are also some power lines which have no GEOTEM response whatsoever.

The power line monitor, of course, is of great assistance in identifying cultural anomalies of this type. It is important to note, however, that geological conductors in the vicinity of power lines may exhibit a weak response on the monitor because of current induction via the earth.

Fences, pipelines, communication lines, railways and other man-made conductors can give rise to GEOTEM responses, the strength of which will depend on the arounding of these objects.

Another facet of this analysis is the line-to-line comparison of anomaly character along suspected man-made conductors. In general, the amplitude, the rate of decay, and the anomaly width should not vary a great deal along any one conductor, except for the change in amplitude related to terrain clearance variation. A marked departure from the average response character along any given feature gives rise to the possibility of a second conductor.

In most cases a visual examination of the site will suffice to verify the presence of a man-made conductor. If a second conductor is suspected the ground check is more difficult to accomplish. The object would be to determine if there is (i) a change in the man-made construction, (ii) a difference in the grounding conditions, (iii) a second cultural source, or (iv) if there is, indeed, a geological conductor in addition to the known man-made source.

### B. Surficial Conductors

This term is used for geological conductors in the overburden, either glacial or residual in origin, and in the weathered layer of the bedrock. Most surficial conductors are probably caused by clay minerals. In some environments the presence of salts will contribute to the conductivity. Other possible electrolytic conductors are residual soils, swamps, brackish ground water and alluvium such as lake or river-bottom deposits, flood plains and estuaries.

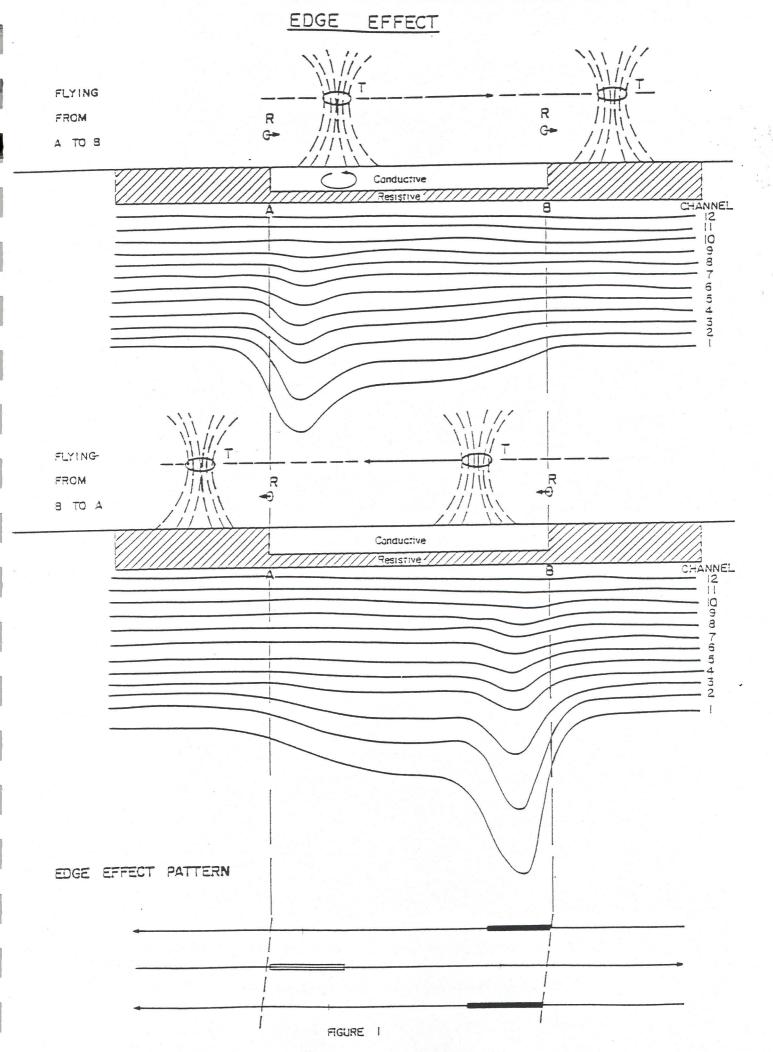
Normally, most surficial materials have low to intermediate conductivity so they are not easily mistaken for highly conductive bedrock features. Also, many of them are wide and their anomaly shapes are typical of broad horizontal sheets.

When surficial conductivity is high it is usually still possible to distinguish between a horizontal plate (more likely to be surficial material) and a vertical body (more likely to be a bedrock source) thanks to the asymmetry of the GEOTEM responses observed at the edges of a broad conductor when flying adjacent lines in opposite directions. The configuration of the system is such that the response recorded at the leading edge is more pronounced than that registered at the trailing edge. Figure 1 illustrates the "edge effect" and the resulting conductive pattern in plan view. In practice there are many variations on this very diagnostic phenomenon.

One of the more ambiguous situations as to the true source of the response is when surface conductivity is related to bedrock lithology as for example, surface alteration of an underlying bedrock unit. At times, it is also difficult to distinguish between a weak conductor within the bedrock (e.g. near-massive sulphides) and a surficial source.

In the search for massive sulphides or other bedrock targets, surficial conductivity is generally considered as interference but there are situations where the interpretation of surficial-type conductors is the primary goal. When soils, weathered or altered products are conductive, and in-situ, the GEOTEM responses are a very useful aid to geologic mapping. Shears and faults are often identified by weak, usually narrow, anomalies.

Analysis of surficial conductivity can be used in the exploration for such features as lignite deposits, kimberlites, paleochannels and ground water. In coastal or arid areas, surficial responses may serve to define the limits of fresh, brackish and salty water.



### C. Bedrock Conductors

This category comprises those anomalies which cannot be classified according to the criteria established for cultural and surficial responses. It is difficult to assign a universal set of values which typify bedrock conductivity because any individual zone or anomaly might exhibit some, but not all, of these values and still be a bedrock conductor. The following criteria are considered indicative of a bedrock conductor:

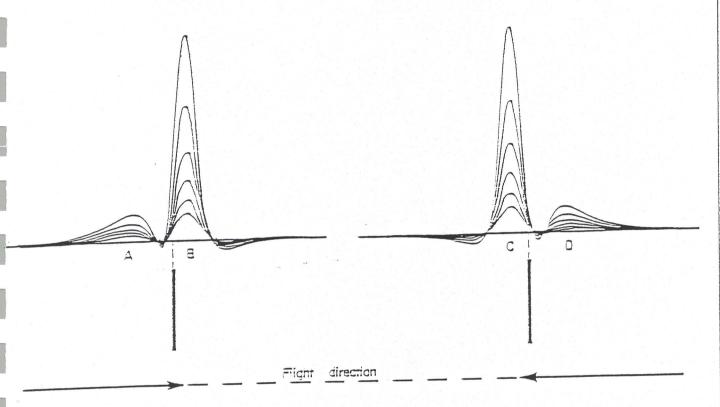
- An intermediate to high conductivity identified by a response with slow decay, with deflections most often present in the later channels.
- 2. The anomaly should be narrow, relatively symmetrical, with a well-defined peak.
- 3. There should be no serious displacement of anomaly position or change in anomaly shape (other than mirror image) with respect to flight direction, except in the case of non-vertical dipping bodies. The alternating character of the response as a result of line direction can be diagnostic of conductor geometry. Figures 2 to 6 illustrate anomalies associated with different target models.
- 4. A small to intermediate amplitude. Large amplitudes are normally associated with surficial conductors. The amplitude varies according to the depth of the source.
- 5. A degree of continuity of the EM characteristics across several lines.
- 6. An associated magnetic response of similar dimensions. One should note, however, that those rocks which weather to produce a conductive upper layer will possess this magnetic association. In the absence of one or more of the characteristics defined in 1, 2, 3 and 4, the related magnetic response cannot be considered significant.

Most obvious bedrock conductors occur in long, relatively monotonous, sometimes multiple zones following formational strike. Graphitic material is usually the most probable source. Massive syngenetic sulphides extending for many kilometres are known in nature but, in general, they are not common. Long formational structures associated with a strong magnetic expression may be indicative of banded iron formations.

A bedrock conductor reflecting the presence of a <u>massive sulphide</u> would normally exhibit the following characteristics:

- a high conductivity,
- a good anomaly shape (narrow and well-defined peak),
- a small to intermediate amplitude,
- an isolated setting,

# THE VERTICAL PLATE RESPONSE



ANOMALY MAP PRESENTATION (no lag applied)

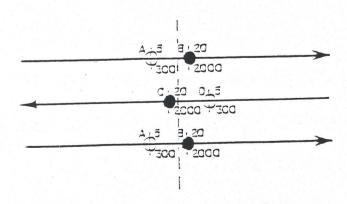


Figure 2

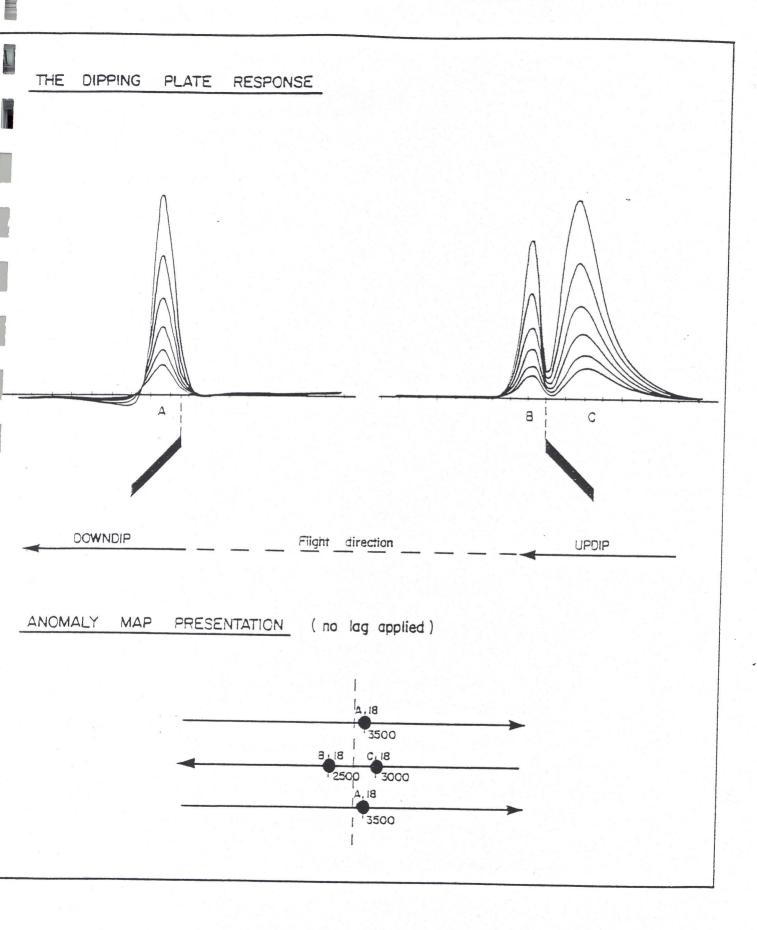
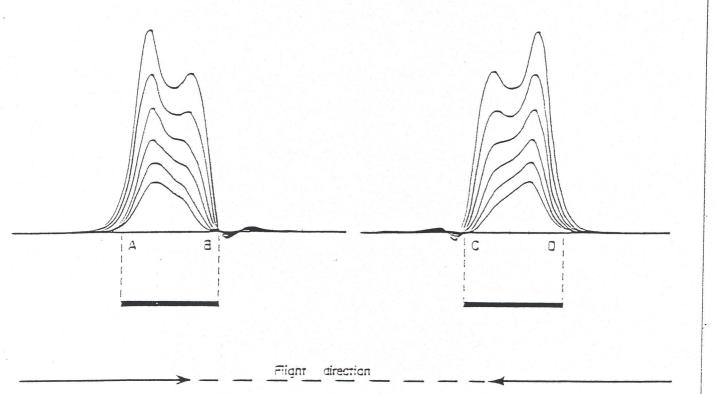


Figure 3

# THE HORIZONTAL PLATE RESPONSE



ANOMALY MAP PRESENTATION ( no log applied )

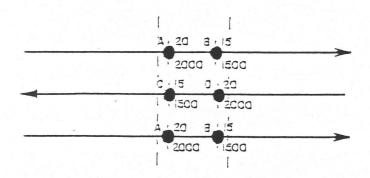
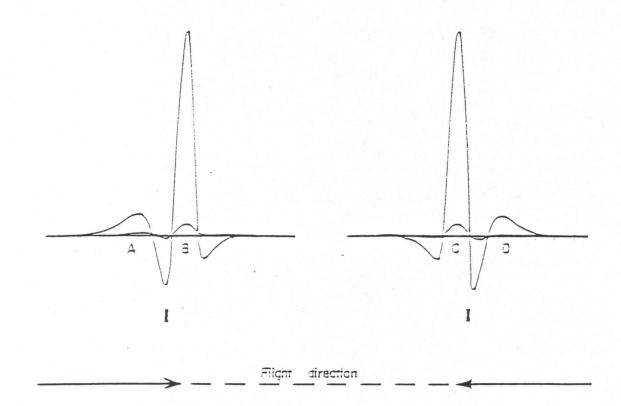


Figure 4

# THE VERTICAL RIBBON RESPONSE



ANOMALY MAP PRESENTATION (no lag applied)

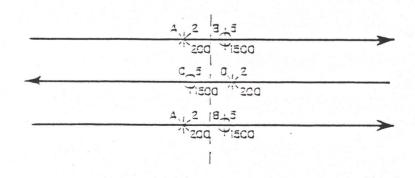


Figure 5

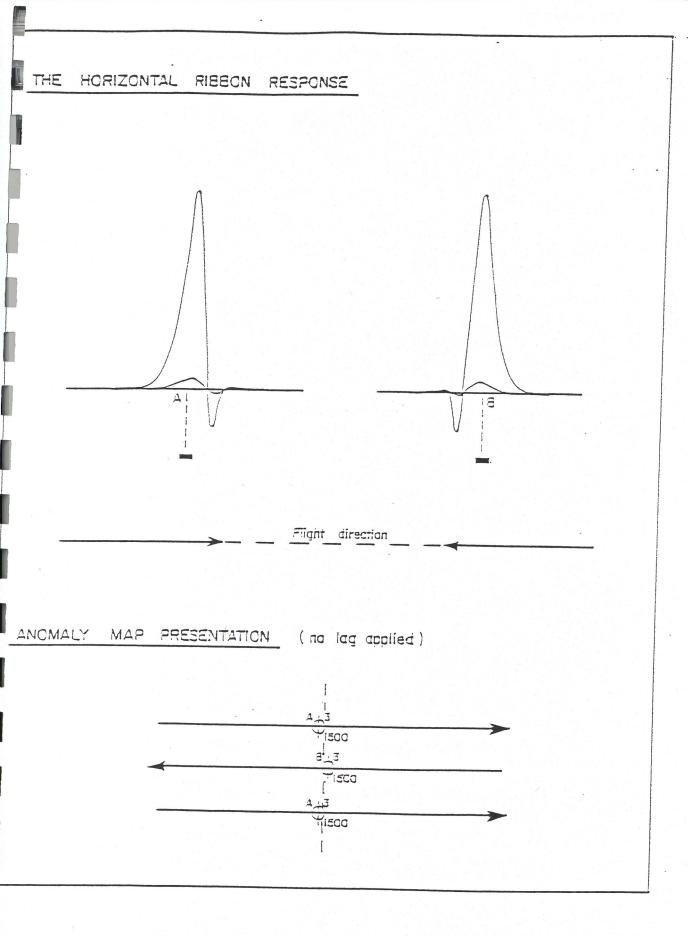


Figure 6

- a short strike length (in general, not exceeding one kilometre), and
- preferably, with a localized magnetic anomaly of matching dimensions.

The selection of targets from within extensive (formational) belts is much more difficult than in the case of isolated conductors. Local variations in the EM characteristics, such as in the amplitude, decay, shape etc., can be used as evidence for a relatively localized occurrence. Changes in the character of the EM responses, however, may be simply reflecting differences in the conductive formations themselves rather than indicating the presence of massive sulphides and, for this reason, the degree of confidence is reduced.

Another useful guide for identifying localized variations within formational conductors is to examine the magnetic data compiled as isomagnetic contours. Further study of the magnetic data can reveal the presence of faults, contacts and other features which, in turn, help define areas of potential economic interest.

Finally, once ground investigations begin, it must be remembered that the continual comparison of ground knowledge to the airborne information is an essential step in maximizing the usefulness of the GEOTEM data.

# Appendix D The Utility of Multicomponent Time-Domain Electromagnetic Measurements

# Theutility of multicomponent time-domain electromagnetic measurements

GEOTERREX
OTTAWA, CANADA
1994

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### ABSTRACT

Time-domain airborne electromagnetic systems historically measure the in-line horizontal (x) component. New versions of the hardware are designed to collect two additional components (the vertical (z) and horizontal lateral (y) components) to provide greater diagnostic information.

In areas where the geology is near horizontal, the z-component response provides greater signal to noise, particularly at late delay times. This allows the conductivity to be determined to greater depth. In a layered environment, the symmetry implies that the y component will be zero; hence a non-zero y component will indicate a lateral inhomogeneity.

Three components can be combined to give the energy envelope of the response. Over a vertical plate, this function has a single positive peak and no side lobes. The shape of the energy envelope is dependent on the flight direction, but less so than the x component.

In the interpretation of discrete conductors, the z-component data can be used to ascertain the dip and depth to the conductor using simple rules of thumb. When the profile line is perpendicular to the strike direction and over the center of the conductor, the y component will be zero; otherwise it will be a combination of the x and z components. The extent of contamination by the x and z components can be used to ascertain the strike direction and the lateral offset of the target respectively.

Having the z- and y-component data increases the total response when the profile line has not traversed the target. This increases the possibility of detecting a target located between adjacent flight lines or beyond a survey boundary.

### INTRODUCTION

The acquisition of multiple-component electromagnetic (EM) data is becoming more commonplace. In some techniques, such as those which use the plane-wave assumption (MT, CSAMT and VLF) more than one component has been acquired as a matter of routine for some time (see reviews by Vozoff, 1990, 1991; Zonge and Hughes, 1991; McNeill and Labson, 1991). Other controlled-waveform finite-source systems which have acquired multiple component data are generally experimental, such as those described in the appendices of Spies and Frischknecht (1991).

Slingram EM systems, comprising a moving dipolar transmitter and a moving receiver, generally only measure one component of the response. The MaxMin system was designed to have the capability to measure a second (minimum coupled) component, but this does not appear to be used in practice. The only systems which have two receiver coils are those which measure the wavetilt or polarization ellipse (Frishknecht et al., 1991).

Historically, time-domain EM systems have been capable of collecting multicomponent data in a sequential manner by reorienting the sensor for each component direction. The usefulness of additional components is discussed by Macnae (1984) for the case of the UTEM system. Macnae concluded that as extra time was required to acquire the additional components, this time was better spent collecting more densely spaced vertical-component data. These data, which are less subject to spheric noise, could subsequently be converted to the horizontal components via the Hilbert transform operator because the measured response is a potential field. Multiple-component data was routinely collected by Newmont Exploration using their proprietary EMP system (no longer in use). Examples describing the interpretation of EMP data are given by Barnett (1984) and Boyd and Wiles (1984).

Recent instrument developments have been towards multicomponent systems. For example, commercially available ground systems such as the Geonics PROTEM,

the Zonge GDP-32 and the SIROTEM have been expanded to include multiple input channels which allow three (or more) components to be acquired simultaneously. There is also a version of the UTEM system with multiple channels currently being developed at Lamontagne Geophysics Ltd. The multichannel receivers require complimentary multicomponent sensors — for ground based systems these have been developed by Geonics Ltd and Zonge Engineering and Research Organization.

Multicomponent borehole measurements are performed most efficiently with multicomponent borehole sensor probes. Prototype probes were developed at Biliden (Hodges et al., 1991) and by Lee (1986), but multicomponent measurements are now becoming more common with the release of probes for commercial use by Crone Geophysics and Exploration Ltd and Geonics. Three component UTEM and SIROTEM borehole sensors are also in development at Lamontagne and Monash University, respectively. Hodges et al. (1991) present an excellent discussion of techniques which can be used to interpret three-component borehole data.

Airborne systems such as frequency-domain helicopter electromagnetic methods acquire data using multiple sensors, but each receiver has a corresponding transmitter in a position of minimum or maximum coupling (Palacky and West, 1991); hence, these systems are essentially multiple single-component systems. The exception to this rule is the now superseded Dighem III system (Fraser, 1972) which used one transmitter and three receivers.

The only multicomponent airborne EM (AEM) system currently in operation is the SPECTREM system (Macnae, et al., 1991). This system, a proprietary development of the PROSPECT system (Annan, 1986), was originally designed to acquire the x, y and z components, but is apparently only collecting two components (x and z) at the time of writing. Other multicomponent systems currently in development are:

- i) the SALTMAP system,
- ii) a helicopter time-domain system (Hogg, 1986), and
- iii) a new version of the GEOTEM® system (GEOTEM is a registered

trademark of Geoterrex).

Apart from a few type curves in Hogg (1986), there is little literature available which describes how to interpret data from these systems.

This paper is intended to give an insight into the types of responses expected with the new multicomponent AEM systems, and the information which can be extracted from the data. The insight could be of some assistance in interpreting data from multicomponent moving-source ground EM systems, should this type of data be acquired, but would not necessarily be applicable to fixed-source EM.

The use of multicomponent data will be discussed for a number of different applications. For illustration purposes, this paper will use the transmitter-receiver geometry of the GEOTEM system (Figure 1), which is comparable to the other fixed-wing geometries (SPECTREM and SALTMAP). The GEOTEM system is a digital transient EM system utilizing a bipolar half-sinusoidal current waveform [more details are in Annan and Lockwood (1991)]. The sign convention used in this paper is shown in Figure 1, with the y component being into the page.

### SOUNDING IN LAYERED ENVIRONMENTS

In a layered environment, current flow is horizontal (Morrison et al., 1969) so the z component is much larger than the x component, particularly in resistive ground and/or at late delay times. At the same time, the spheric noise in the vertical (z) direction is five to ten times less (Macnae, 1984; McCracken et al., 1986), so the z component has a greater signal-to-noise ratio. Figure 2 shows theoretical curves over two different, but similar, layered earth models. One model is a half-space of 500  $\Omega$ m and the other is a 350 m thick layer of 500  $\Omega$ m overlying a highly resistive basement. In this plot the data have been normalized by the total primary field. The z component is six to ten times larger than the x component, and both curves are above the noise level, at least for part of the measured transient. For this plot, the spheric noise was assumed not to be contaminating the x component, so a 30 ppm

noise level would be typical. In order to distinguish between half-space and 350 m thick layer, it is not only necessary that each curve be above the noise level, the difference between the response of the two models must also be greater than the noise level. Figure 3, which is this difference, shows that only the z-component difference is above the noise level and hence able to distinguish the two conductivity models. Hence for the case shown, the z component is more useful than the x component for determining whether there is a resistive layer at 350 m depth. Because of the z-component signal is generally larger in a layered environment, levels this conclusion will also be true in general.

In practice, the receiver bird changes orientation, so part of the vertical component will be measured by the horizontal sensors. This could be corrected using positioning information and/or redundant information from the primary field. An alternative is to calculate and model the magnitude of the total field, as this is a quantity independent of the receiver orientation. This is what was done by Macnae et al. (1991) when calculating the conductivity depth sections for SPECTREM data.

The symmetry of the secondary field is such that if the EM bird is aligned radially with the transmitter, the y component will always be zero. In fact, the y component will be zero whenever the conductivity structure on both sides of the aircraft is the same. A non-zero y component is therefore useful in identifying off-line lateral inhomogeneities in the ground.

# DISCRETE CONDUCTORS

Historically, airborne transient electromagnetic (TEM) data have been used for conductor detection. One of the reasons why the old INPUT system measured the x component of the response was that the response measured peaked when the receiver passed over the top of the conductor. The bottom part of Figure 4 shows the response over a vertical conductor, as calculated with the PLATE program (Dyck and West, 1984). As is standard practice, the data have been plotted at the receiver, so the x component peaks over the conductor. Note that there is also a peak at 200 m, just

before the transmitter passes over the conductor, and a trailing edge negative to the left of the conductor. The z component (dashed line) has two peaks and a large negative trough just before the conductor. Because of the symmetry, the y component (dotted line) is zero. All the peaks, troughs and negatives make the response of a single conductor complicated to display and hence interpret. The display can be simplified by calculating the "energy envelope" of the response. This quantity is the square root of the sum of the squares of each component and its spatial derivative. As such, it is similar to the analytic signal in magnetics, but as it is not "analytic" in the formal mathematical sense, it is termed the energy envelope. The energy envelope plotted on Figure 4 is almost symmetric, and would be a good quantity to present, in plan form, as contours, or as an image.

What little asymmetry remains in the energy envelope is a good indication of the coupling of the AEM system to the conductor and can be used to reduce the asymmetry of the individual curves. For example, the top part of Figure 4 shows the x and z component normalized by the energy envelope at each point. The size of the two x peaks and the two z peaks are now roughly comparable.

# Dip determination

The response of a plate with a dip of  $120^{\circ}$  is shown on Figure 5. For the x and z components, the peak on the down dip side is larger. However, the z-component peaks remain identifiable over a larger range of dips and plotting the ratio of the peak sizes, as has been done with solid squares on Figure 6, shows that the ratio is very close to the tangent of the dip divided by 2. Hence, calculating the ratio of the peak amplitudes (R) will yield the dip  $(\theta)$  using

$$\theta = 2 \tan^{-1}(R)$$
.

# Depth determination

As the depth of the body increases, the distance between the two z-component peaks increases. As an example of this, Figure 7 shows the case of a plate 150 m

deeper than the plate of Figure 4. The peaks are now 450 m apart, as compared with 275 m on Figure 4. A plot of the peak-to-peak distances for a range of depths is shown on Figure 8 for plates with 60, 90 and 120° dips. The distribution follows a straight line and it can be seen that if the dip is within 30 degrees of vertical, the depth to the top of the body can be determined to within about 25 m, an error which is tolerable in airborne EM interpretation. More traditional methods for determining the depth (Palacky and West, 1973) use the rate of decay of the response in all the channels. Our method uses only one channel, allowing a depth for each delay time to be ascertained and, hence, it may be possible to track any migration of the current system in the conductor.

# Strike and offset determination

The response shown in Figure 4 varies in cases when the plate has a strike different from 90° or the flight path is offset from the center of the plate.

Figure 9 shows the response for a plate with zero offset and Figure 10 is when the plate is offset by 150 m from the profile line. The z and x response are little changed from the no offset case, but the y response, which was zero, is now significant. In fact, the shape of the response curve appears to be the mirror image of the z response.

In the case when the plate strikes at 45°, the y component is similar in shape but opposite in sign to the x-component response (Figure 11).

These similarities can be better understood by looking at schematic diagrams of the secondary field from the plate. Figure 12 shows a plate and the field in section, with increasing offset denoted by moving the aircraft to the right. For zero offset the field is vertical (z only). As the offset increases, the field rolls over into the y component.

The case of variable strike is depicted in plan view on Figure 13 by leaving the plate stationary and changing the flight direction. As the strike changes from perpendicular, what was purely x component now rotates into the y component.

The y component  $(H_y)$  can thus be considered to a be a mixture of x  $(H_z)$  and z  $(H_z)$  components, viz

$$H_{y} = C_{sik} H_{z} + C_{off} H_{z},$$

an equation which is approximate only. The response for a variety of strike angles and offset distances has been calculated and in each case the y-component response has been decomposed into the x and z components by solving for the constants of proportionality  $C_{stk}$  and  $C_{off}$ .

A plot of  $C_{stk}$  for the case of zero offset and varying strike direction  $\zeta$  is seen on Figure 14. The values determined from the data are plotted with solid squares and compared with the tan of 90° minus the strike angle. Because the agreement is so good, the formula

$$\zeta = 90 - \tan^{-1}(C_{stk})$$

can be used to determine the strike. This relation was first obtained by Fraser (1972).

When the strike is fixed at 90°, and the offset varies, the corresponding values obtained for  $C_{off}$  have been plotted with solid squares on Figure 15. Again, there is good agreement with the arctangent of  $C_{off}$  and the  $\zeta$  angle between vertical and the line joining the profile line with the center of the plate. If the depth d is already obtained using the method described above, or by the method described in Palacky and West (1973), then the offset distance O can be obtained from the formula

$$O = d \tan (\zeta)$$
$$= d C_{\text{off}}$$

# Lateral detectability

Figure 12 illustrates that the y component becomes relatively strong as the lateral displacement from the conductor is increased. Thus, if the y component is measured then the total signal will remain above the noise level at larger lateral displacements from the conductor. This has been illustrated by assuming a flat-lying conductor, here approximated by a wire loop circuit of radius 125 m (Figure 16). The x, y and z components of the response have been computed using the formulae for the

large-loop magnetic fields in Wait (1982). The results are plotted on Figure 16 as a function of increasing lateral displacement of the transmitter/receiver from the center of the conductor. The transmitter and receiver are separated in a direction perpendicular to the lateral displacement, so as to simulate the case when the system is maximal coupled to the conductor, but the flight line misses the target by an increasing amount. The effect of varying the conductance or measurement time has been removed by normalizing the response to the total response measured when the system is at zero displacement. At displacements greater than 80 m, the y component is clearly larger than any other component. If the same sensitivity and noise level is assumed for the x, y and z components, it is clearly an advantage to measure the y component to increase the chances of detecting the target when the flight line has not passed directly over the conductor.

### CONCLUSIONS

The additional components (y and z) collected with multicomponent AEM data improve the interpretability of the data.

The z-component data enhances the ability of the AEM system to resolve layered structures. This is because of the reduced spheric noise and the increased signal. If all the components are employed to correct for coil rotation, then the data quality and resolving power is increased further, as individual components are not contaminated by another component. Having better signal-to-noise and greater fidelity in the data will allow deeper layers to be interpreted with confidence.

The y component is helpful in detecting when the conductivity structure has lateral inhomogeneities which are not symmetric about the flight line.

All components are used to calculate the energy envelope, which is a valuable quantity to image. The energy envelope has a single peak over a vertical conductor and two peaks over a dipping conductor, one at either end. The asymmetry in the response profiles for the individual components can be reduced by normalizing each

profile by the energy envelope.

All three components are of great use in determining the characteristics of discrete conductors. For example, the distance between the two peaks in the z-component profile can be employed to determine the depth. Also, the normalized ratio of the two z-component peaks helps to ascertain the dip of the conductor. The x component has been used in the past for these purposes, but is not as accurate, as it requires the data at all delay times, or an ability to identify a very small peak.

The y component can be utilized to extract information about the conductor which cannot be obtained from single component AEM data. The degree of mixing between the y and z components can give the lateral offset of the conductor (provided the depth is known), while the mixing between the y and x component gives the strike of a vertical conductor.

Finally, measuring the y component gives an enhanced ability to detect a conductor away from the profile line, either between lines or beyond the edge of a survey boundary.

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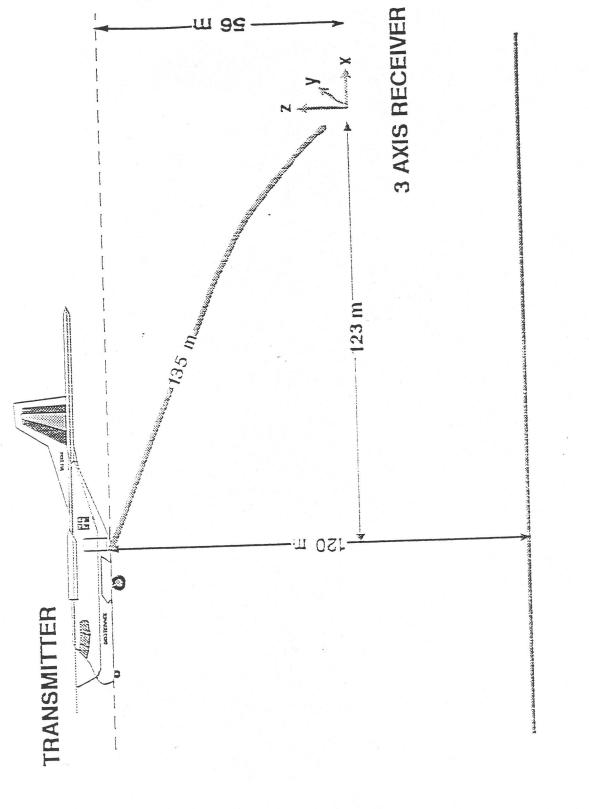
### FIGURE CAPTIONS

- Fig. 1. The geometric configuration of the GEOTEM® system. The system comprises a transmitter on the aircraft and a receiver sensor in a "bird" towed behind the aircraft. The z direction is positive up, x is positive behind the aircraft and y forms a right-hand coordinate system.
- Fig. 2. The response for a 500 Ω·m half-space (solid line) and a 500 Ω·m layer of thickness 350 m, overlying a resistive half-space (dashed line). The zcomponent responses are the two curves with the larger amplitudes and the two x-component response curves are six to ten times smaller than the corresponding z component. A noise level of 30 ppm is considered to be a typical in the absence of strong spherics.
- Fig. 3. The difference between the z and x components of the response for the two models of Figure 2. Only the z-component difference is above the noise level for any time and is, therefore, the only component capable of distinguishing between the responses of the two models.
- Fig. 4. (Bottom) The response of a 600 by 300 m plate 120 m below an aircraft flying from right to left. The plotting point for the response is below the receiver. The x-component response is the smaller amplitude solid line, the z component is the dashed line and the y-component response is the dotted line. The larger amplitude solid line is the energy envelope of all three components. (Top) The z and x components normalized by the energy envelope. These curves are for a delay time of 0.4 milliseconds after the transmitter current is turned off.
- Fig. 5. (Bottom) same as Figure 4, except the plate is now dipping at 120°. On the top graph note that the down-dip (left) peak on the normalized z-component response is now larger that the right peak.
- Fig. 6. The ratio of the peak amplitudes of the normalized z-component response (left/right), plotted with solid squares. The ratio plots very close to the tangent

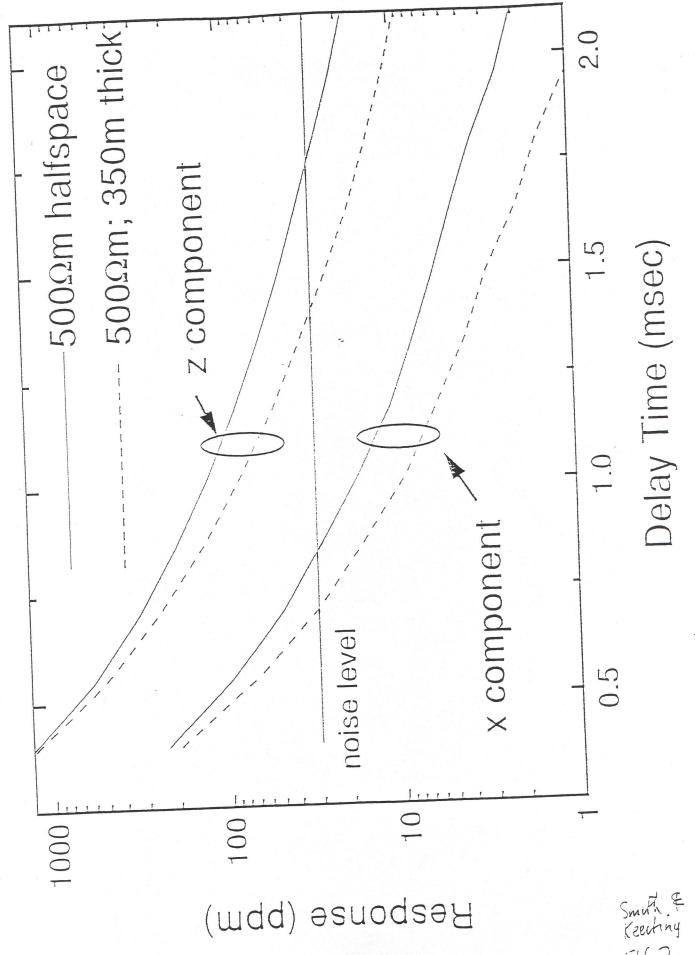
- of half the dip angle of the plate.
- Fig. 7. The same as Figure 4, except the plate is now 270 m below the aircraft. Note that the distance between the z-component peaks is now much greater.
- Fig. 8. The peak-to-peak distance as a function of plate depth for three different dip angles. A variation in dip of ± 30° does not result in a large change in the peak-to-peak distance.
- Fig. 9. The response of a 300 by 300 m plate traversed by a profile line crossing the center of the plate in a direction perpendicular to the plate (strike angle of the plate with respect to the profile line is 90°).
- Fig. 10. Same as Figure 9, except the profile line has been offset from the center of the plate by -150 m in the y direction (equivalent to a 150m displacement of the plate).
- Fig. 11. Same as Figure 9, except the profile line traverses the plate such that the strike angle of the plate, with respect to the profile line, is 45°.
- Fig. 12. A schematic diagram of the plate and the magnetic flux of the secondary field (section view). For increasing offset from the center of the plate, the magnetic field at the receiver rolls over from the z to the y component.
- Fig. 13. A schematic diagram of the plate and the magnetic flux of the secondary field (plan view). Here varying strike is depicted by an equivalent variation of the flight direction. As the flight direction rotates from a strike angle of 90°, the receiver rotates so as to measure a greater response in the y direction.
- Fig. 14. The ratio of the components y to x  $(C_{nk})$ , plotted as a function of varying strike angle (solid squares). The data agree very closely with the cotangent of the strike angle.
- Fig. 15. The arctangent of the ratio of the components y to z  $(C_{off})$ , plotted as a function of varying offset (solid squares). There is good agreement with the angle of the profile line from the center of the plate (measured in degrees from the vertical).
- Fig. 16. Plan view of a grey circular flat-lying conductor, radius 125 m. The AEM

system is offset a distance d from the center of the conductor along a radial line where the transmitter and receiver couple strongly to the body. The traverse direction of the system is from the bottom to the top of the figure.

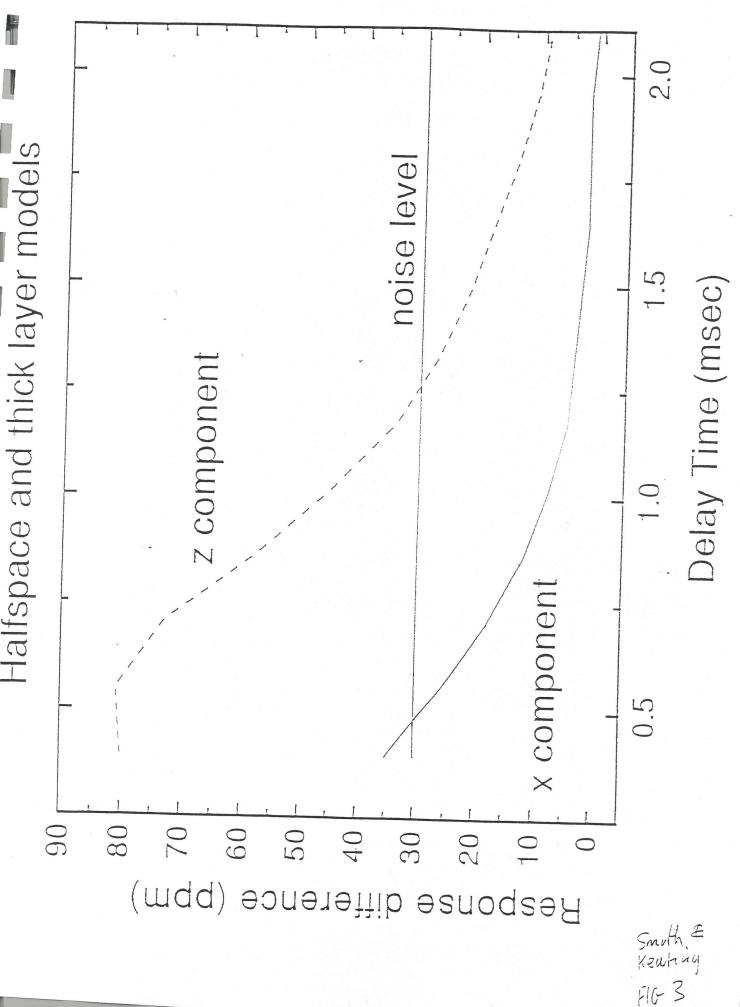
Fig. 17. The normalized response of the EM system plotted as a function of increasing offset distance. The x component falls off most rapidly and the y component most slowly with offset distance.

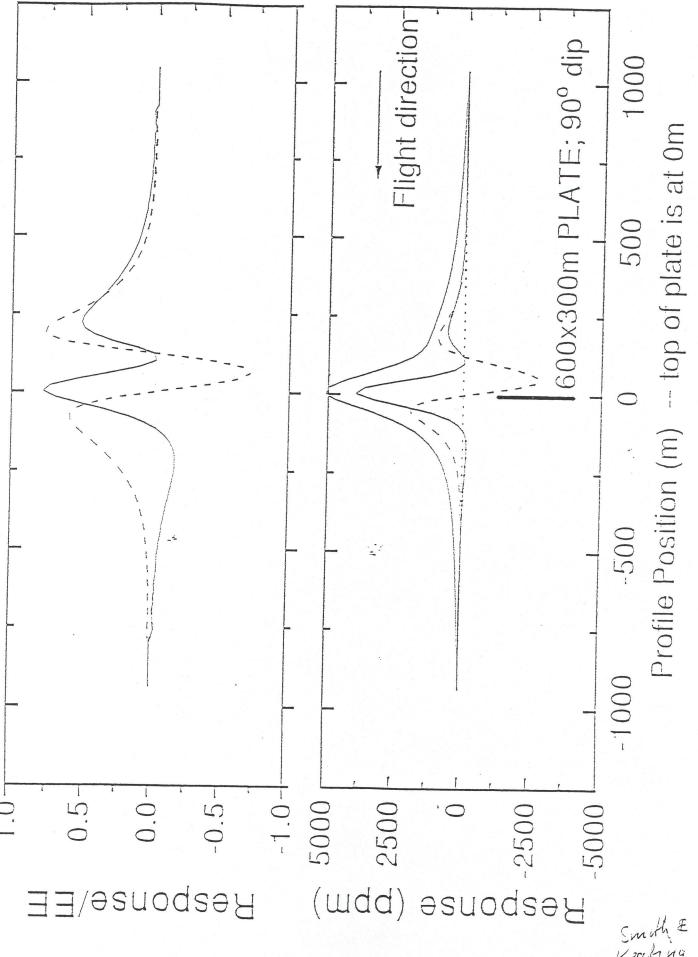


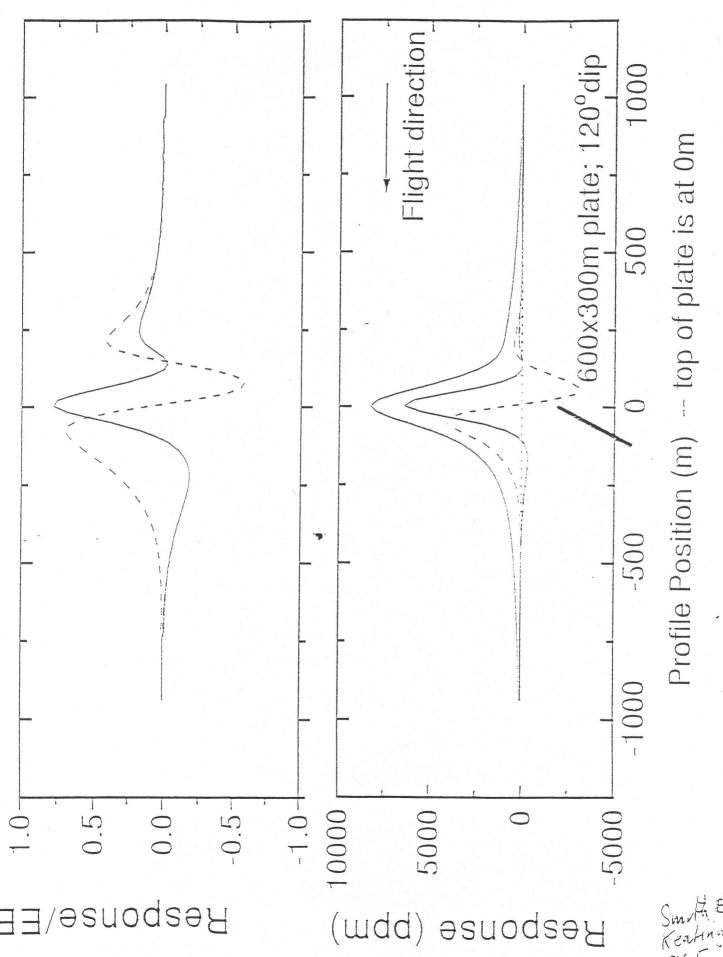
Smith & Keating

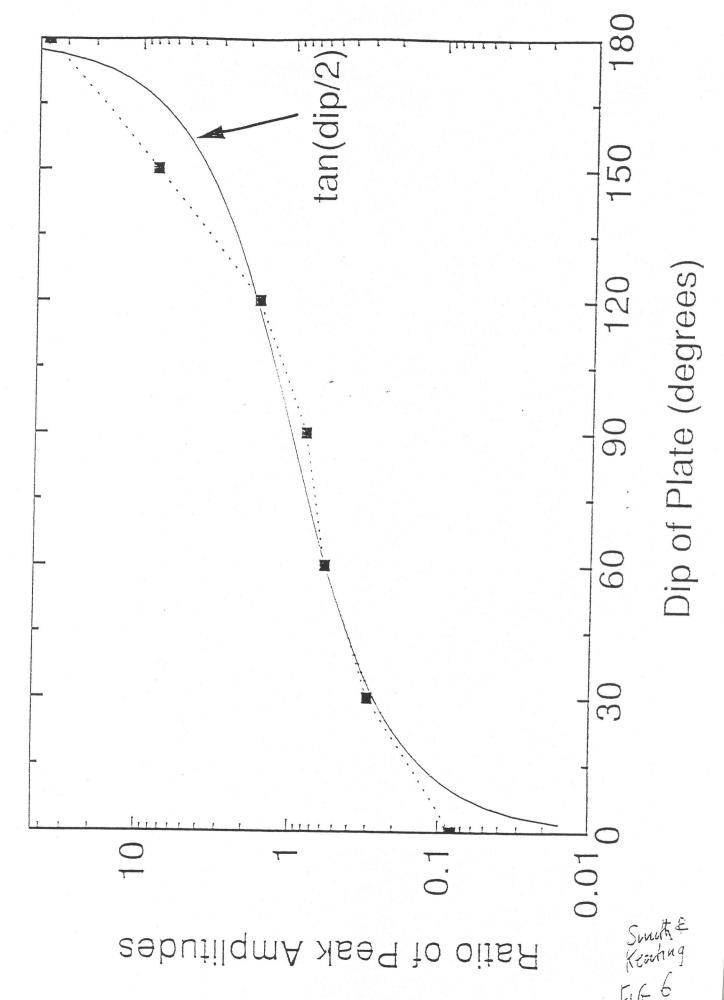


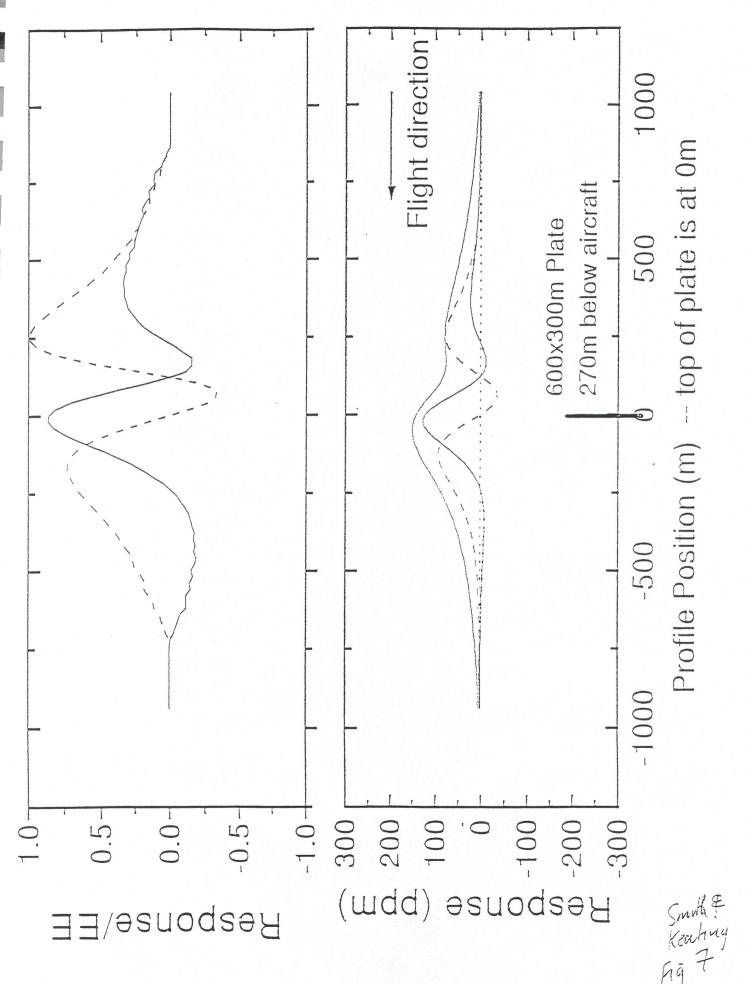
F16-2

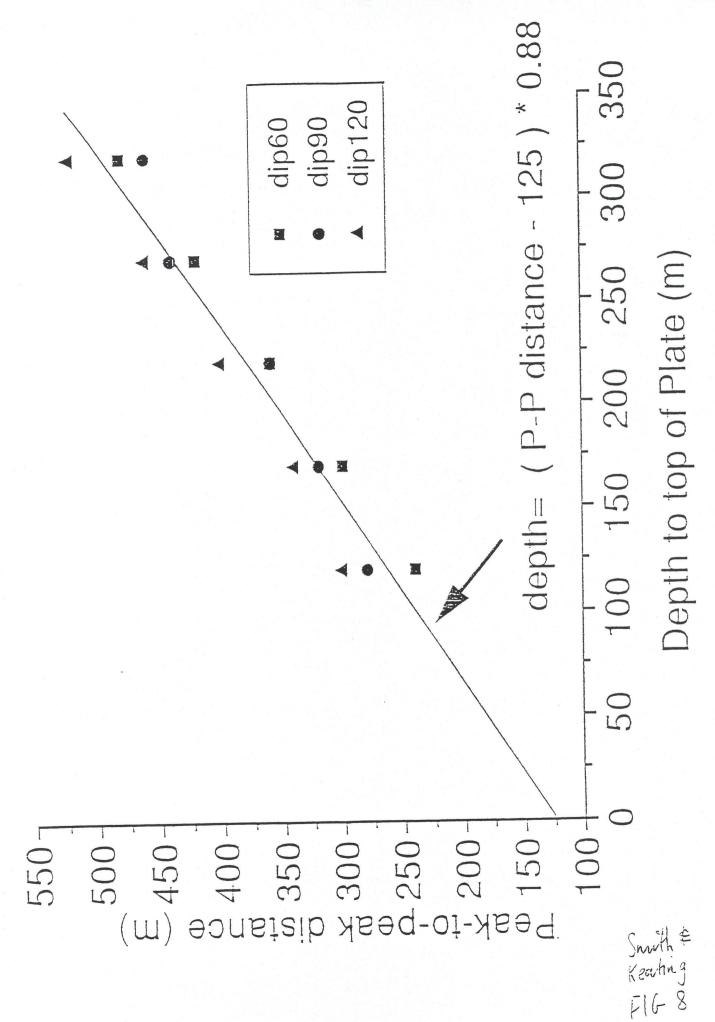


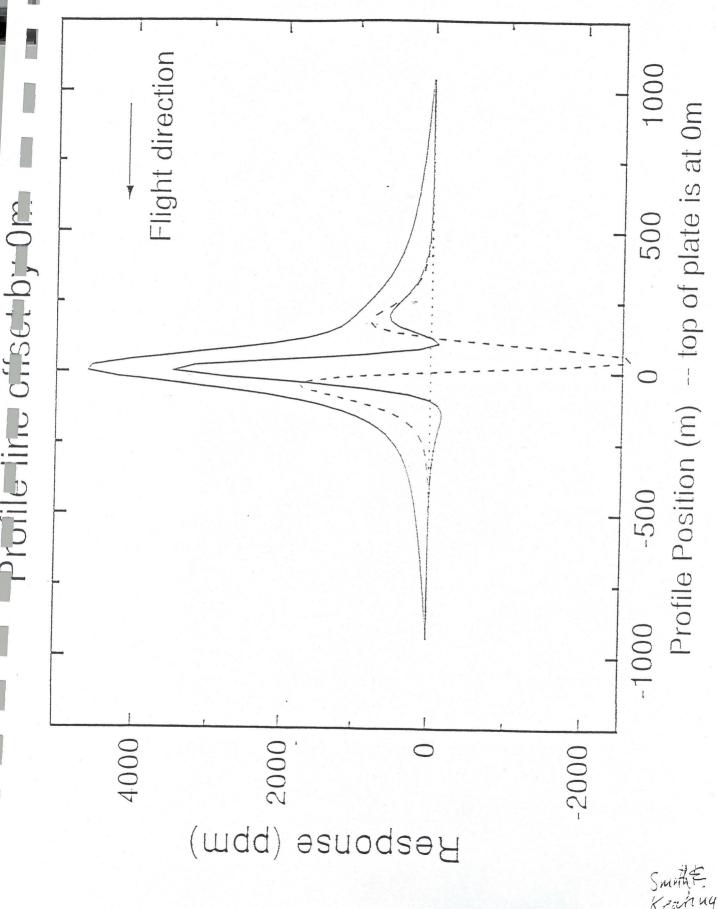




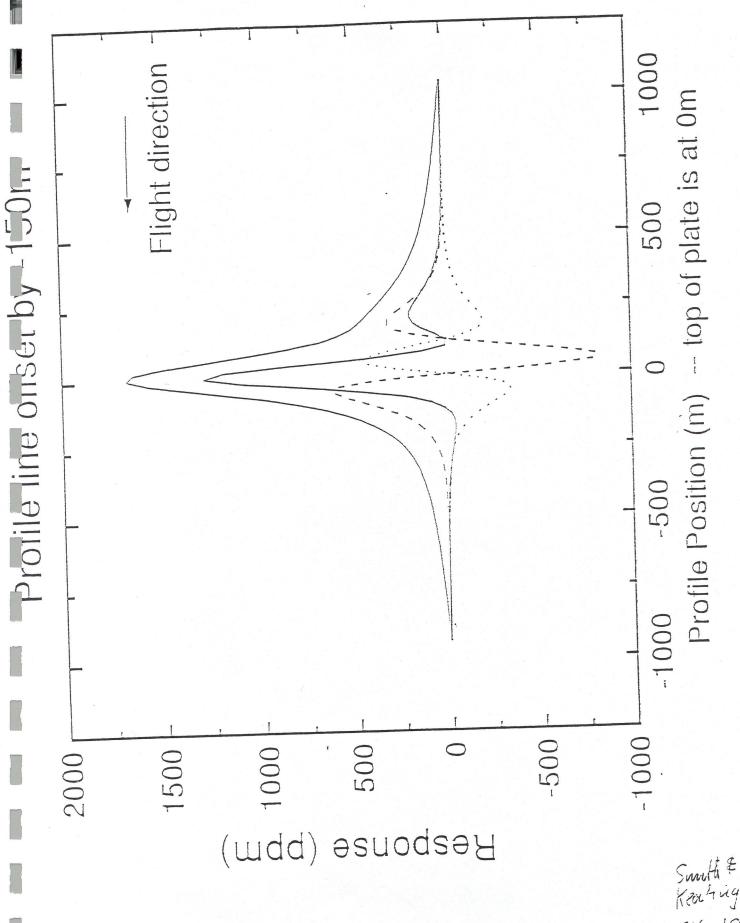




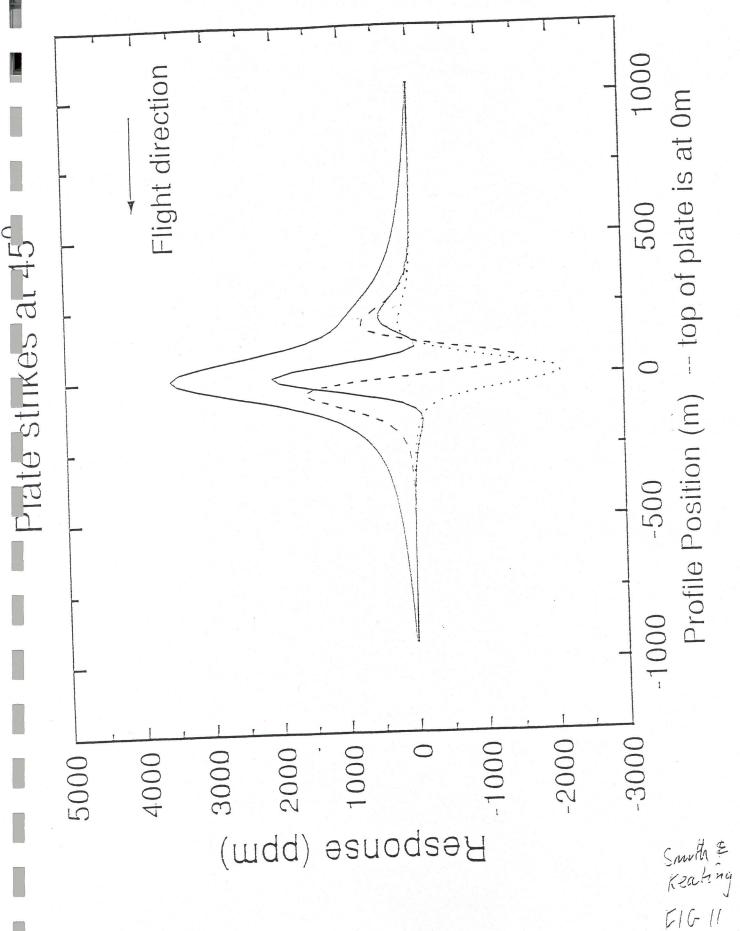


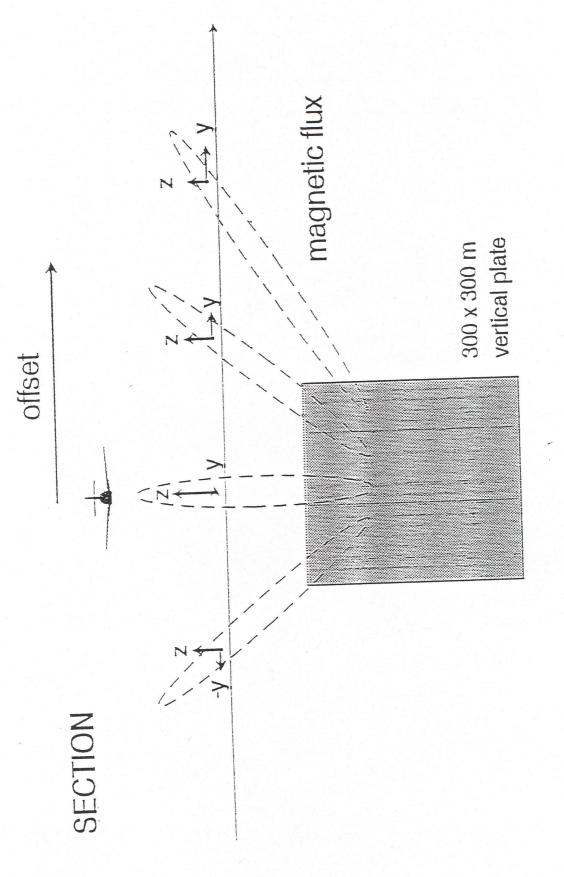


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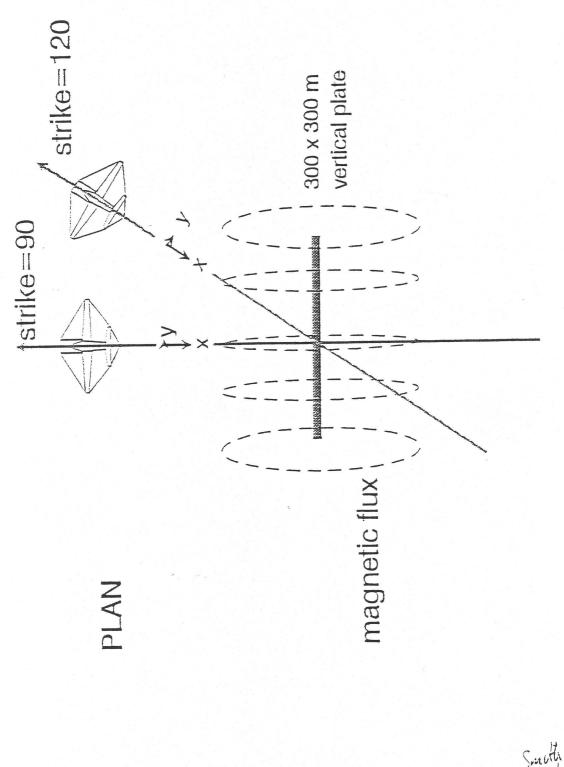


Smith & Kentaig PIG-10

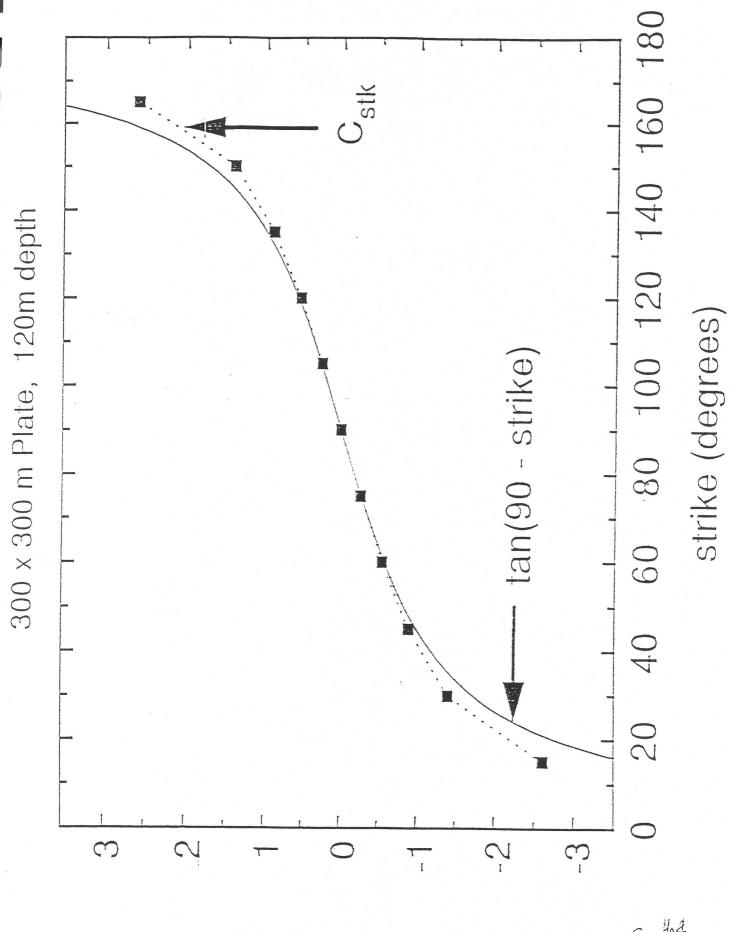




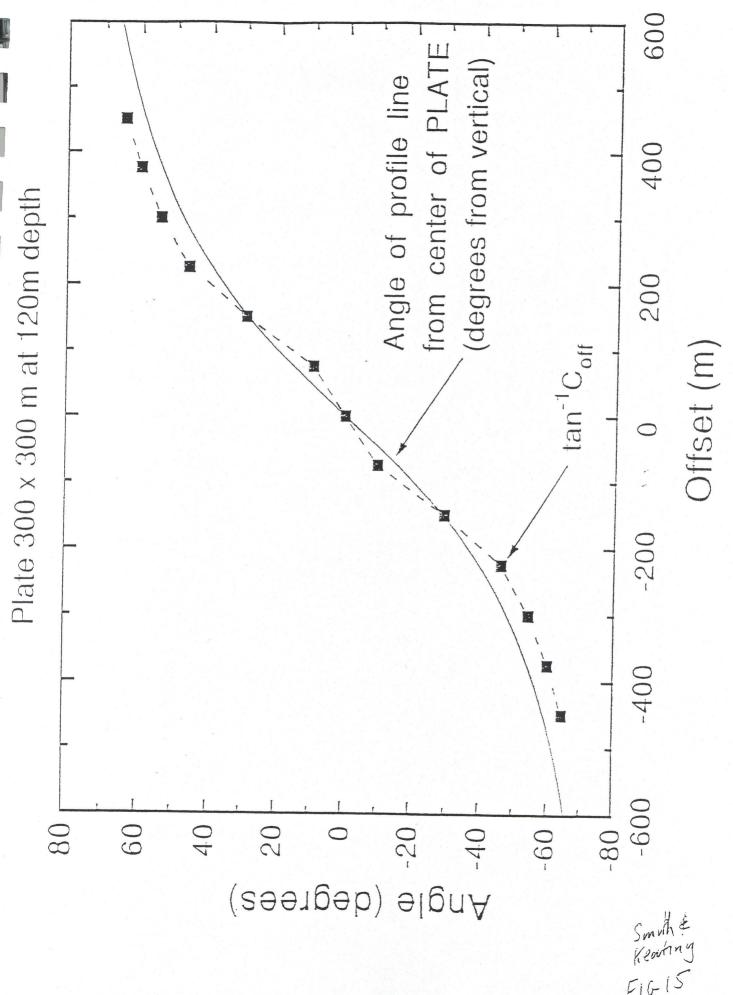
Sunda & Keahing FIG-12

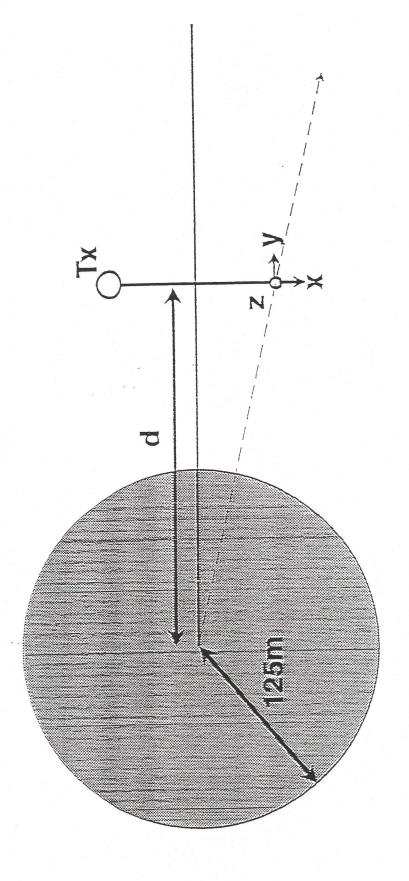


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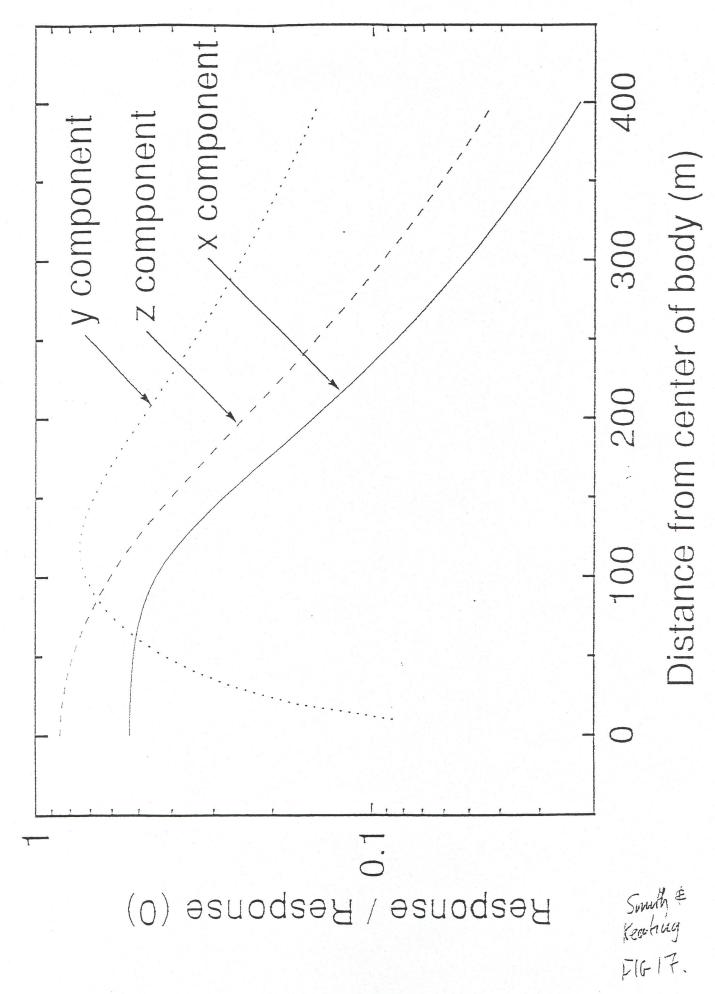


Smuth & Keating Fig 15





Smith & Kealing



# Appendix E Multicomponent GEOTEM® Modelling

## Multicomponent GEOTEM modelling

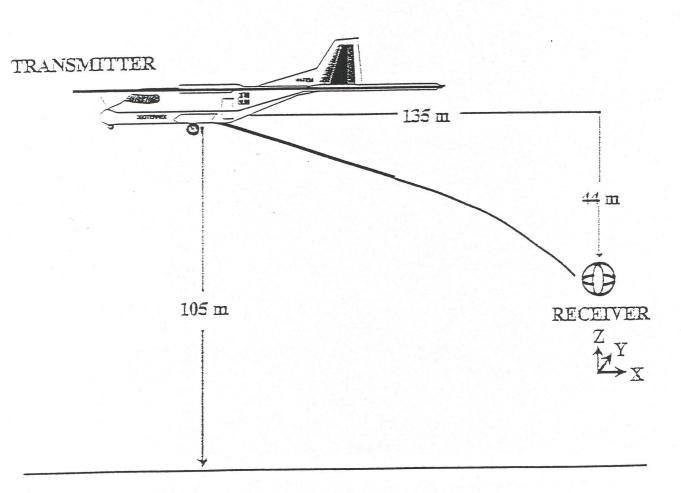
The PLATE program has been used to generate synthetic responses over a number of plate models with varying depth of burial (0, 150 and 300m) and dips (0, 45, 90 and 135 degrees). The geometry assumed for the GEOTEM system is shown on the following page, and the transmitter waveform on the subsequent page. For simplicity, only six receiver gates have been calculated and plotted.

In all cases the plate has a strike length of 600 m, with a strike direction into the page. The width of the plate is 300 m. As the flight path traverses the centre of the plate, the y component is zero and has not been plotted.

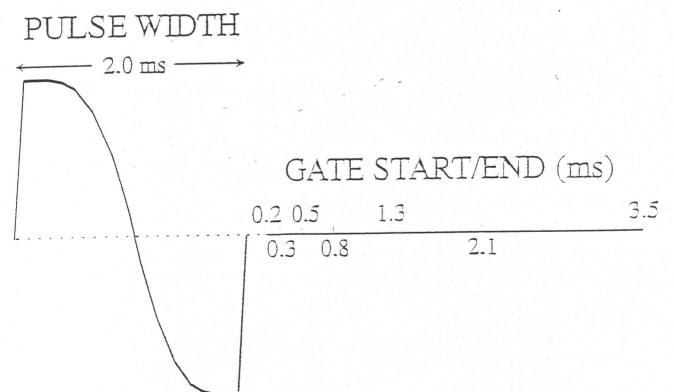
The conductance of the plate is 20 S. In cases when the conductance is different, an indication of how the amplitudes may vary can be obtained from the nomogram included.

In the following plots all components are normalized to the total primary field.

# GEOTEM Geometry for modelling



# Transmitter Waveform and Receiver sampling



# Nomogram

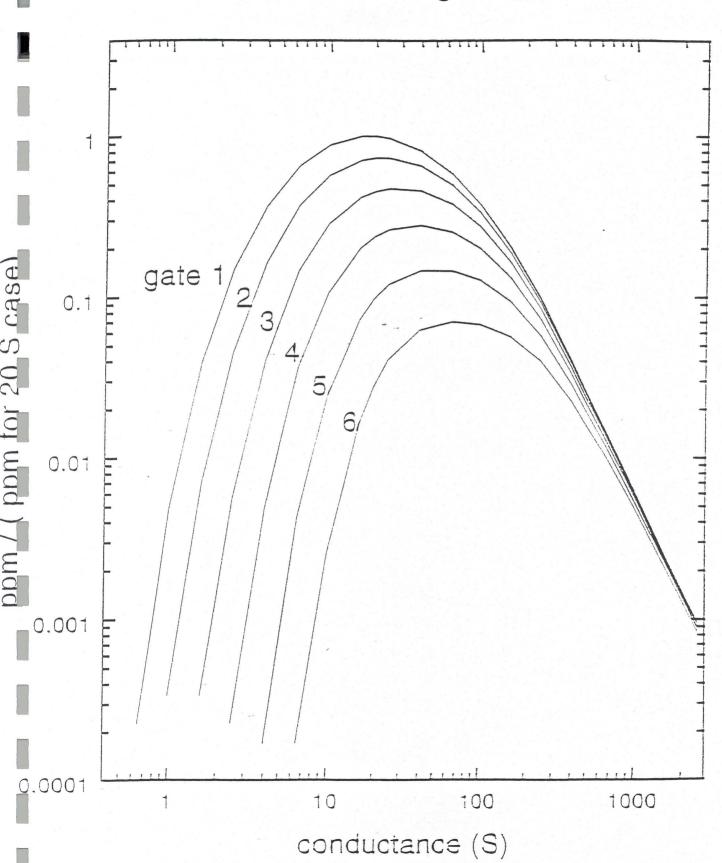


Plate: depth =0; dip =0

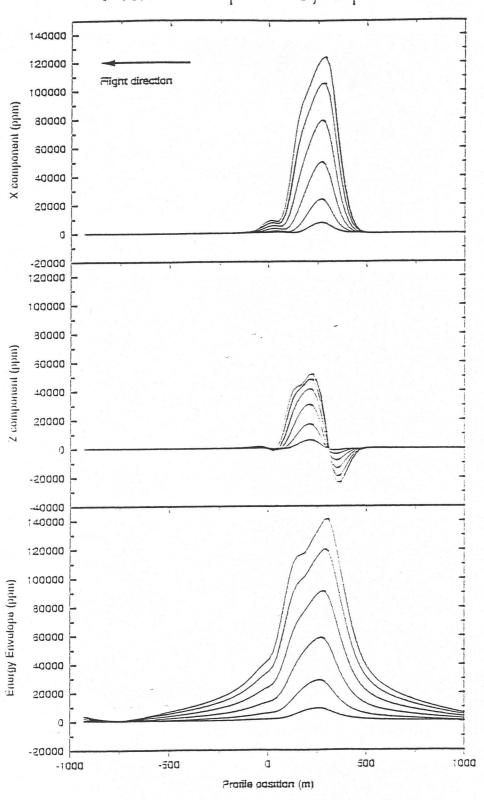


Plate: depth =150; dip =0

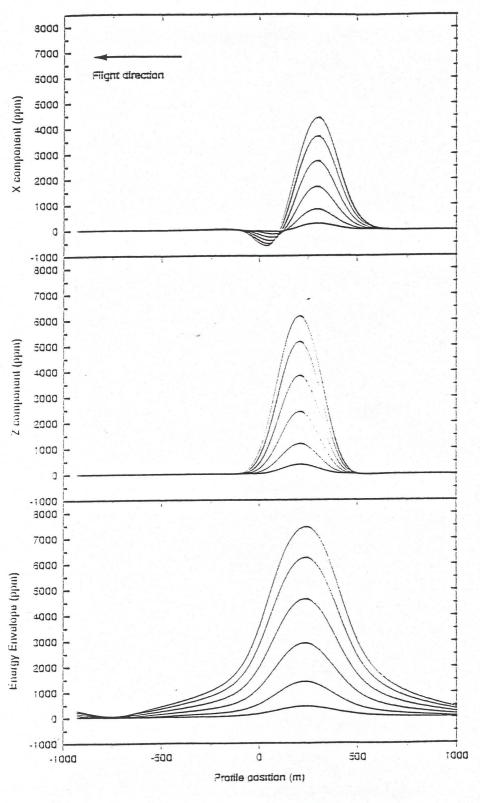


Plate: depth =300; dip =0

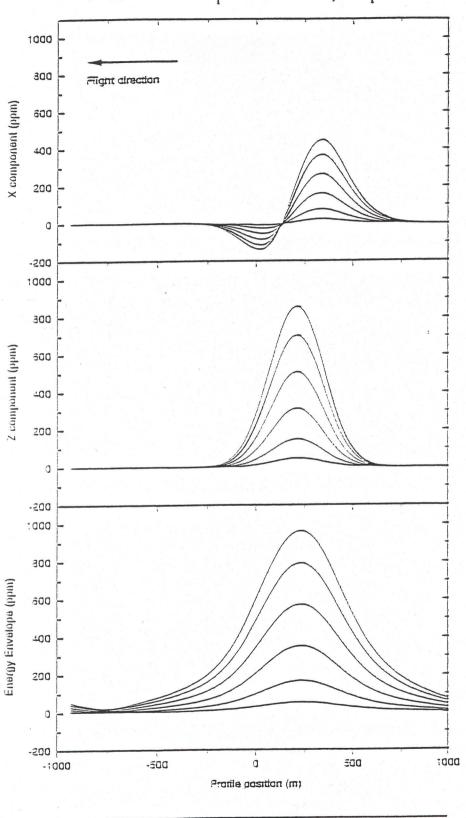


Plate: depth =0; dip =45

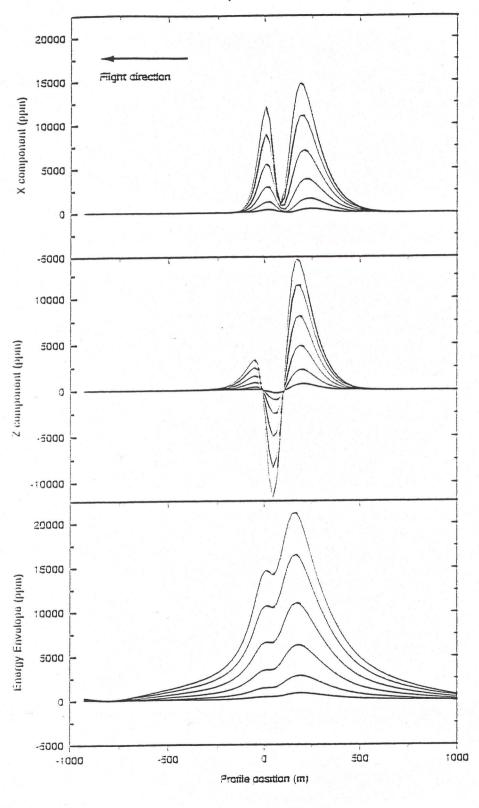


Plate: depth =150; dip =45

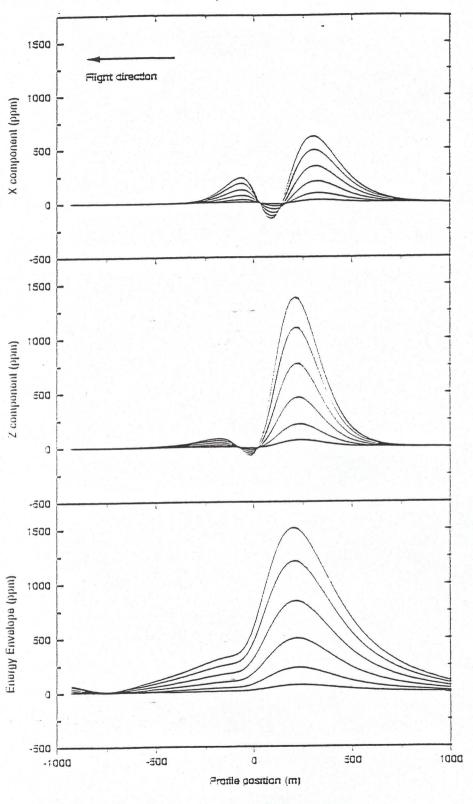


Plate: depth =300; dip =45

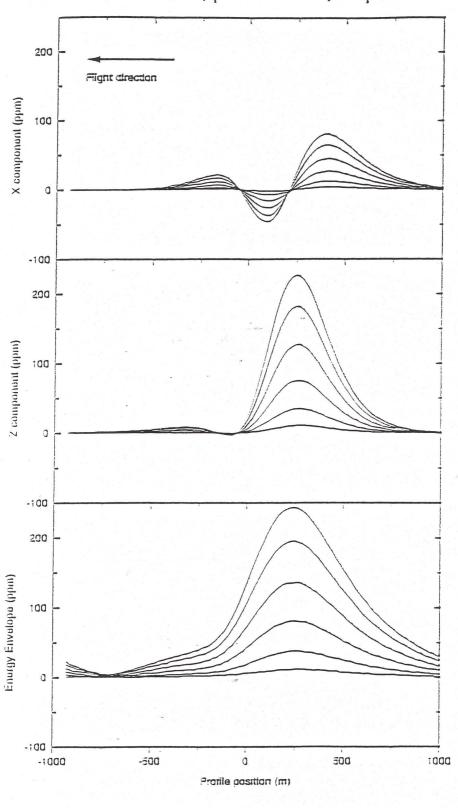


Plate: depth =0; dip =90

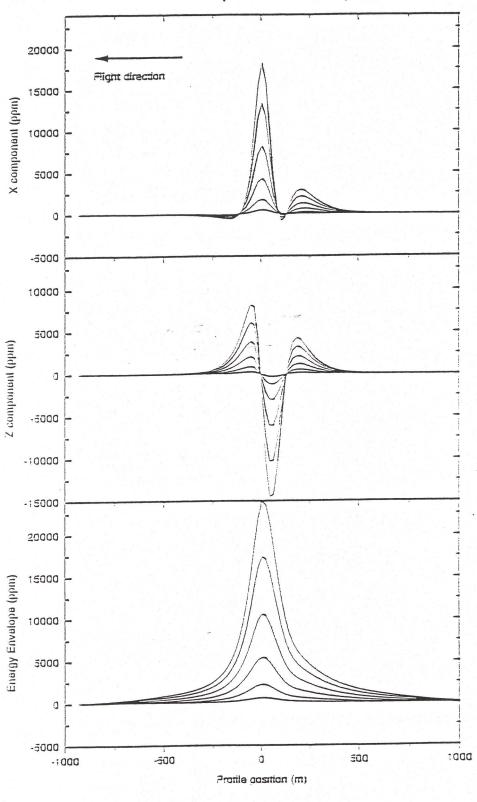
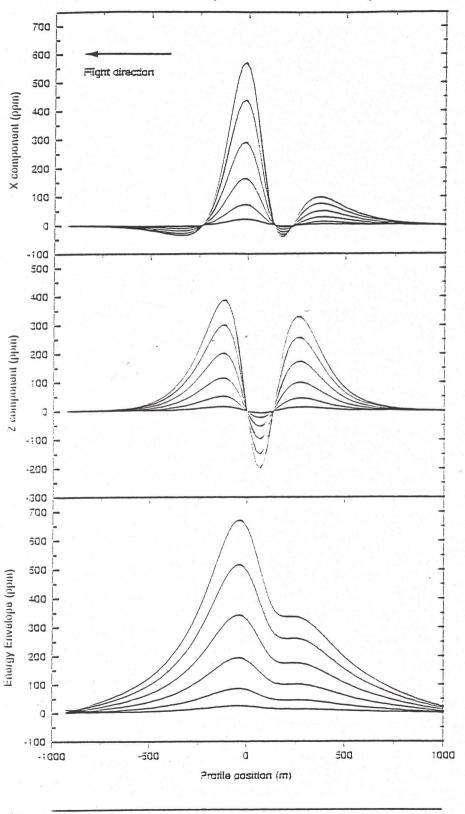


Plate: depth =150; dip =90



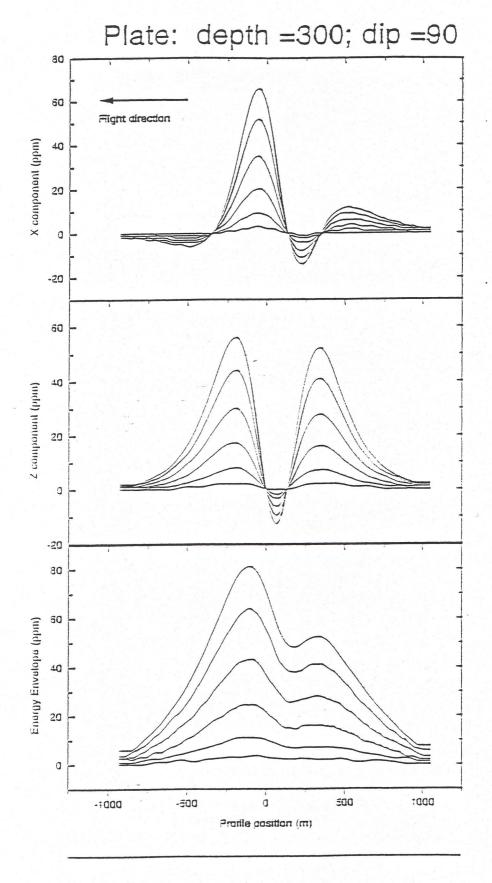


Plate: depth =0; dip =135

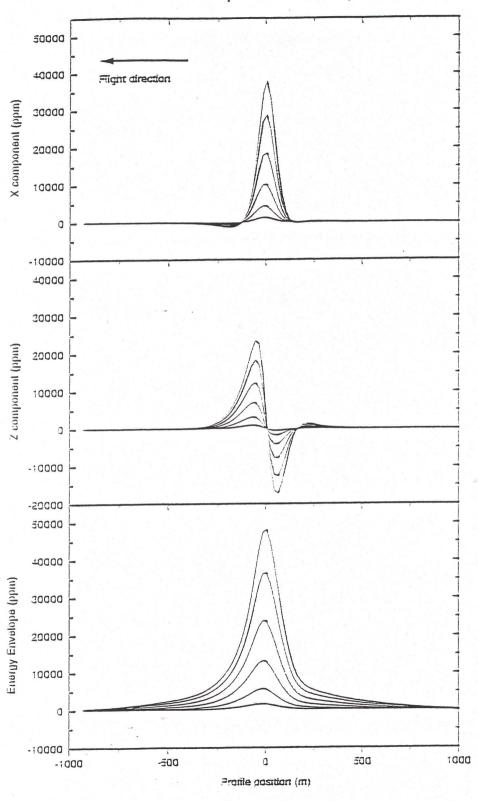


Plate: depth =150; dip =135

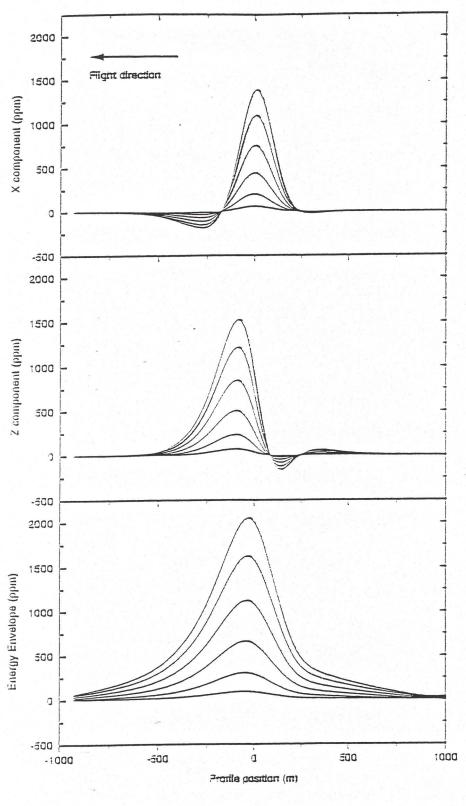
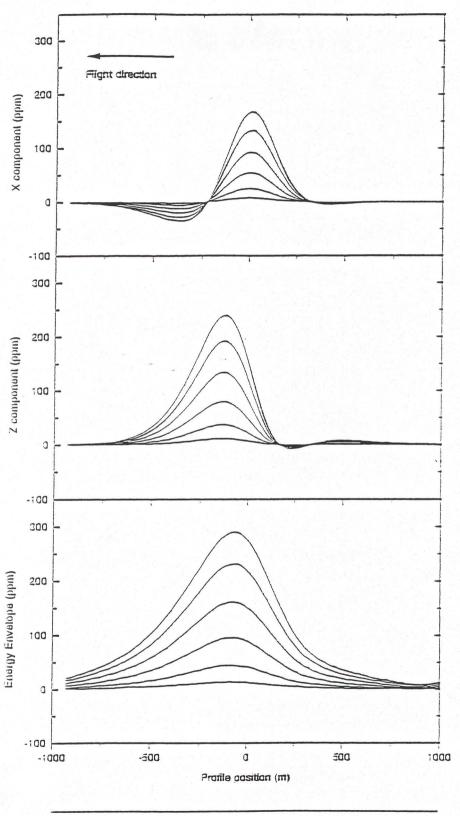


Plate: depth =300; dip =135



# Appendix F Archive Format Description

### GPS ACCURACY CHECK

Flown at the Sierra Vista airport March 14,1997

File: gpscheck.001

ASCII coded 13 I 10 Record length = 130 + CR = 131

One record per bloc

### **FORMAT**

CHARACTE	RS VARIABLE	UNITS	
1-10 11-20 21-30 31-40 41-50 51-60 61-70 71-80 81-90 91-100 101-110 111-120 121-130	Line identifier Fiducial GPS Latitude (WGS84) GPS Longitude (WGS84) GPS Latitude (Clarke 1866) GPS Longitude (Clarke 1866) UTM Easting (Clarke 1866) UTM Northing (Clarke 1866) GPS Elevation GPS Acquired time GPS UTC Time GPS Satellite designator GPS PDOP	Line x 100 + part Seconds DD.dddddd DD.dddddd DD.dddddd DD.dddddd DD.dddddd metres	
131-131	CR		

Data is archived as one sample per second.

LINE-ID	DIR FST-FID LST-FID TA	PE BLOCK	REC SAMPLES

 <sup>1
 0
 52362
 60430
 1
 1
 1
 8069

 2
 0
 61045
 62380
 1
 8070
 1
 1336</sup> 

\*\*\*\*\* ARCFLT COMPLETED \*\*\*\*\*

#### MAG & EM LAG CHECK

Flown near the Sierra Vista airport March 14,1997

File: magemlag.001

ASCII coded 5 I 10 Record length = 50 + CR = 51 One record per bloc

### **FORMAT**

CHARACTERS	VARIABLE	UNITS
1-10	Line identifier	Line x 100 + part
11-20	Fiducial	Seconds
21-30	Raw x-coil channel 11	ppm
31-40	Raw Magnetics	nT x 100
41-50	Distance	metres
51-51	CR	

Data is archived as ten samples per second.

## LINE-ID DIR FST-FID LST-FID TAPE BLOCK REC SAMPLES

101	3	59708	59748	1	1	1	401	
201	1	59838	59878	1	402	1	401	
301	3	59974	60024	1	803	1	501	
401	1	60115	60155	1	1304	1	401	

\* \* \* \* \* \* ARCFLT COMPLETED \* \* \* \* \* \* \*

#### ALTIMETER CALIBRATION CHECK

Flown at the Sierra Vista airport March 14,1997

File: altcal.001

ASCII coded 7 I 10 Record length = 70 + CR = 71 One record per bloc

#### **FORMAT**

CHARACTERS	VARIABLE	UNITS
1-10	Line identifier	Line x 100 + part
11-20	Fiducial	Seconds
21-30	UTM Easting (Clarke 1866)	metres
31-40	UTM Northing (Clarke 1866)	metres
41-50	Radar altimeter	feet
51-60	Barometric altimeter	feet
61-70	GPS Elevation	metres x 100
71-71	CR	

Data is archived as one sample per second.

LINE-ID	DIR FST-FID LST-FID TAPE	BLOCK	REC SAMPLES

1 0 66722 68081 1 1 1 1360

\* \* \* \* \* \* ARCFLT COMPLETED \* \* \* \* \* \* \*

# Appendix G Contents of CDT grids

## CONTENT OF THE CONDUCTIVITY-DEPTH-TRANSFORM GRIDS

X values = Distance along the flight line.

The grid spacing will vary slightly from one line to the next reflecting the change in the average flying speed along each line.

The number of grid cells along the line is defined by the length of the line in metres, devided by the scale and multiplied by 1000, and this value rounded up to the nearest multiple of 64.

eg. For a line which has a length of 25,500 metres, and created for a scale of 1/50,000, the number of grid cells will be:

and 510/64 = 7.969 rounded up to  $6 \times 64 = 512$ 

The actual grid spacing in the X direction will then be the length of the line divided by the number of gric cells determined.

eg. 25,500 / 512 = 49.805 m

Y values = Dept values.

For the desired maximum depth of investigation, 128 rows are allocated, resulting in a fixed Y direction cell size equal to the desired maximum depth divided by 128.

eg. 600 m / 128 = 4.688 m

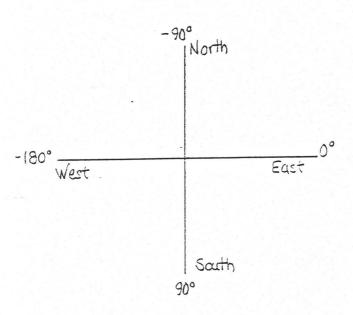
Z values =  $\frac{\text{Computed resistivity values}}{\text{These values are actually stored as (log_{10} of conductivity in Siemens/m) multiplied by 1000.$ 

The conversion to resistivity in ohm-m is then:

,Antilog<sub>10</sub>(value,1000)<sup>-1</sup>

eg.  $-2500 = (Antilog_{10} - 2.5)^{-1} = 316$  ohm-m

Grid rotation angle: This angle is the heading of the flight line based on a Cartesian notation, and therefore will change by 180° for reciprocal line directions.



### Example from line 619 of the Cochiti Block:

Length of line is 27.767 Km

Maximum cepth of investigation is 1300 metres below the reference datum, selected to be 2500 metres above sea level.

Horizontal scale 1/100,000

### Content of the CDT arid file

Number of rows (Y direction)	128
Number of columns ( X direction )	
X grid spacing	
Y grid spacing	
Z multiplier	
Min value	-4.003
May value	0.731

The number of columns and X grid spacing are defined by :

length of line (27767 m)  $\times$  scale (100,000)  $\times$  1000 = 277.67

then. 277.67/64 = 4.339 roundec-up to 5

and  $5 \times 64 = 320$  columns

gric spacing is then equal to length of line (27767 m) /320 = 66.772 m

The Y gric spacing is defined by :

Depth of investigation (1300 m) / 128 rows = 10,156 m

The true data range is:

Minimum of  $-4.003 = Antilog_{10}-4.003^{-1} = 10069$  ohm-m

Maximum of  $-0.731 = Antilog_{10} - 0.731^{-1} = 5.38$  ohm-m

## Correction of CDT grids for Topography

The CDT grids have moditionally been created with the depth information relative to surface, with surface represented as a horizontal line (top of the CDT grid). In areas of generally flat terrain, this representation is acceptable. However, in more rugged areas, this representation can be misleading and result in false interpretation of the results, if one is not careful; i.e., what may appear to be a steeply dipping layer may actually be a horizontal layer passing under a steep hill, since the surface reference at this point is not horizontal, but rising sharply.

For this reason, the surface reference of the CDT grids has been reprojected to the true terrain surface. The grid file must still however have a horizontal datum reference, which in this case has now been chosen as a datum slightly above the highest barometric elevation reached over the survey black. This is taken into account at the creation of the CDT grid, since the desired maximum depth of investigation must now be specified from the barometric datum chosen and not from surface. This will explain what may at first appear to be an "overly optimistic" depth of investigation.

When viewing the "Terrain corrected" CDT grids, the blank area between the top of the section and the first layer of real values represents air. The contour of the first layer of real values defines the ground surface and should match the true topography (this has been computed by the programme from the radar and barometric altimeter data).

## The fallowing parameters were used for the present survey.

Area	Barometric reference	Max. depth of investigation
Rio Rancho	2200 metres	1100 metres
Rio Puerco	2300 metres	1000 metres
Cochiti	2500 metres	1300 metres

# Appendix H Sample parameter table file

## SAMPLE PARAMETER FILE

125.0000 50.00000 30.00000 130.2083	0000000000 0000000000 0000000000 3333333	= Horiz = Verti = Base = Samp	contal TX-R cal TX-RX Frequency ble Interval	meter table file  X separation in meters  S separation in meters  in Hertz  in micro-seconds  ole number:
20	2	3		
19	4	13		
18	14	23		
17	24	33		
16	34	34		
15	35	35		
14	36	37		
13	38	39		
1	40	41		
2	42	44		
3	45	47		
4	48	51		
5	52	56		
6	57	62		
7	63	69		
8	70	77 86		
9	78 87	98		
10	99	112		
12	113	128		
			veforms in	microVolts:
120	14675	-661		-688
2	911262	-55983	519035	-11901
3	1709430	-97688	925387	-56506
4	1734052	-99747	926706	-104344
5	1678316	-96873	895411	-150269
6	1602132	-92483	854211	-194091
7	1509451	-87169	804378	-235407
8	1401621	-80936	746414	-273788
9	1279878	-73861	681078	-308837
10	1145721	-66110	609107	-340321
11	1000696	-57762	531412	-367955
12	846485	-48785	448744	-391584
13	684608	-39421	362102	-410950
14	516946	-29800	272327	-425877
15	345374	-19827	180536	-436222

16 17 18 19 20 21 22 23 24 25 26 27 28 29 30 31 32 33 34 35 36 37 38 39 40 41 42 43 44 45 46 47 48 49 50 50 50 50 50 50 50 50 50 50 50 50 50	171568 -2335 -174755 -343675 -507725 -664577 -813421 -952399 -1079856 -1195149 -1296663 -1383818 -1455587 -1511438 -1551011 -1573907 -1488026 -261690 -48090 -11557 -11458 -11079 -9577 -8567 -7813 -7099 -6482 -6022 -5575 -5003 -4819 -4439 -4140 -3901 -3676 -3445 -3258 -3009 -2926	-9810 242 10184 20008 29471 38592 47189 55138 62515 69228 75058 80104 84228 87446 89742 91078 86300 12288 4235 1072 819 789 714 634 633 505 520 435 385 448 367 324 330 258 305 274 248 194 232	-97481 -187766 -275230 -359027 -438367 -512352 -580304 -641614 -695566 -741808 -779835 -809328 -830104 -841936 -782325 -113731 -22433 -6349 -6482 -6212 -5394 -4827 -4387 -3937 -3582 -3287 -3056 -2791 -2584 -2439 -2192 -2113 -1949 -1831 -1685 -1603	-380204 -355309 -326896 -295291 -260876 -224051 -185260 -144896 -103391 -61170 -19005 2076 542 475 588 602 615 631 637 654 652 659 659 661 662 664 671 665 674 667 664 670	
52	-3258	248	-1685	664	
54 55	-2926 -2682	232 216	-1494 -1425	661 662	
56	-2620	204	-1355	656	
57 58	-2539 -2296	165 160	-1240 -1207	665 663	
59	-2245	173	-1153	650	
60	-2122	170	-1056	651	

106	-522	68	-186	615
107	-380	69	-224	620
108	-620	4	-190	617
109	-479	84	-181	610
110	-443	27	-206	611
111	-460	49	-169	622
112	-417	34	-187	607
113	-451	57	-153	619
114	-433	77	-164	613
115	-400	44	-132	611
116	412	55	-194	611
117	-381	62	-159	609
118	-399	42	-141	610
119	-361	65	-144	610
120	-342	70	-154	617
121	-349	48	-106	614
122	-363	72	-144	607
123	-344	34	-131	603
124	418	64	-153	607
125	-340	41	-124	609
126	-364	29	-127	608
127	-342	75	-139	613
128	-334	47	-127	609

# CONVERSION FACTORS TO NORMALIZE THE EM DATA FROM PV/M<sup>2</sup> to PPM

FLIGHT N°	X-COIL	Z-COIL
1	0.0248	0.0335
2	0.0359	0.0496
3	0.0350	0.0491
5	0.0354	0.0483
6	0.0337	0.0544
7	0.0352	0.0528
8	0.0352	0.0538
9	0.0347	0.0566
11	0.0324	0.0609
12	0.0346	0.0528
13	0.0344	0.0549
14	0.0329	0.0636
15	0.0320	0.0667
16	0.0326	0.0659
17	0.0326	0.0651
18	0.0318	0.0659

#### Note:

- 1. The conversion factors should be applied as multipliers.
- 2. Flight 1 was a test flight flown with 90 Hz/2 ms, hence the difference between it and the other flights flown with 30 Hz/4 ms.

# Appendix J Production Log

			BASE: ,	ALBUQUERQ		JE, NEW	JE, NEW MEXICO			JOB	JOB NO: 319
Date	ij	Km,	Total	Balance	Hr./Fit.	Total.	REMARKS	Elec.	Oper.	Daily	Km/Hour
			Km.				TO A THE TAX TO A	Operator	Day	Average	
									,	Km.	
Commence of the second			0.0			0.0	VORTION TO THE TOTAL THE TOTAL THE TOTAL TRANSPORT OF THE TRANSPORTION OF THE TRANSPORT THE TRANSPOR	THE CHARLE OF THE ACT OF THE CHARLES AND MANAGEMENT AND THE CHARLES OF THE CHARLES	A THEOREM BROWN WAS	The second secon	A The Comment of the
Feb. 18	ı	0.0	0.0	3,763.0	0.0	0.0	Ground station setup - Pilot flight training.			ERR	ERF
Feb. 19	ı	0.0	0.0	3,763.0	0.0	0.0	Grounded - disabled pilot.		_	0.0	ERR
Feb. 20	1	0.0		3,763.0	0.0	0.0	Pilot flight training.		7	0.0	ERF
Feb. 21	1	0.0	0.0	3,763.0	0.0	0.0	Grounded - awaiting reptacement pilot.		က	0.0	ERR
	_	112.9	112.9	3,650.1	1.5	1.5	Survey test @ 90 Hz - RioRancho.	R.P. + J.M.	4	28.2	75.
	7		169.7	3,593.3	1.6	3.1	Survey test @ 30 Hz - RioRancho.	R.P.	4	42.4	54.7
	3	396.1	565.8	3,197.2	3.0	6.1	Survey @ 30 Hz - RioRancho.	R.P. + J.M.	5	113.2	92.
	1	0.0	565.8	3,197.2	0.0	6.1	Wx out - extremely high winds.		9	94.3	92.
Feb. 25	1	0.0	565.8	3,197.2	0.0	6.1	Wx out - high winds.		7	80.8	92.
	4	0.0	565.8	3,197.2	1.0	7.1	Backup EM bird test flight.	R.P.	80	7.07	79.
Feb. 27	1	0.0	565.8	3,197.2	0.0	7.1	Grounded awaiting flying permit.		6	62.9	79.
	,	0.0	565.8	3,197.2	0.0	7.1	Grounded awaiting flying permit.		10	56.6	79.7
Mar. 1	5	558.8	1,124.6	2,638.4	3.6	10.7	Survey - RioRancho.	R.P. + J.M.	F	102.2	105.1
	စ	485.6	1,610.2	2,152.8	2.9	13.6	Survey - RioRancho.	RP. + J.M.	1	146.4	118.4
Mar. 2	7	428.6	038.	1,724.2	2.9	16.5	Survey - RioRancho.	J.M.	12	169.9	123.
	80		2,326.3	1,436.7	3.0	19.5	Survey - RioRancho.	J.M.	12	193.9	119.3
Mar. 3	တ	0.0	326.	1,436.7	3.8	23.3	Survey - Cochiti. Flight scrubbed - bad mag.	J.M.	13	178.9	.66
	10		2,326.3	436	1.0	24.3	Electronics test flight.	R.P.	13	178.9	95.
	7	63.5	389	1,373.2	1.3	25.6	Survey - Cochiti. Flight aborted - high winds.	R.P.	14	170.7	93.
	12	317.0	907,	1,056.2	3.5	29.1	Survey - Cochiti.	D.P.	15	180.5	93.(
Mar. 6	13	394.8	3,101.6	661.4	3.3	32.4	Survey - Cochiti.	D.P. & J.M.	16	193.9	95.
Mar. 7	14	345.1	446		3.0	35.4	Survey - Cochitl.	D.P. & J.M.	17	202.7	97.
	15	253.9	3,700.6	62.4		38.5	Survey - Cochiti/RioRancho.	D.P. & J.M.	17	217.7	96.1
	16	N	,763.	(0.0)		40.7	Survey - RioPuerco.	D.P.	17	221.4	92.
Mar. 8	17	120.0	3,883.0	(120.0)	2.1	42.8	Survey - additional lines.	D.P. & J.M.	18	215.7	.06
	18	12	995.	(232.0)	4	44.7	Survey - additional lines.	D.P.	19	210.3	89.